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Abstract. The purpose of this work is to study the Etesian winds in the Saroniko Gulf to understand the complicate physical processes that generate the windstorms in the Saroniko Gulf. It is developed a study of the weather patterns characteristic of the Saroniko Gulf and a comparisons of the wind data measured on land by the station of Helliniko airport (years: 1990–2000) and at sea by a buoy (years: 2000–2002) during the month of August. It is presented a statistical model that shows, as a probability of occurrence, the variations of the wind speed in function of the variations of the wind direction. Finally the statistical model is compared with the RAMS limited area model during a test case in August 2004; the now-casting results obtained using these two different methods are compared to data measured by station of the automatic weather network of the HNMS.

1 Introduction

The Etesian winds, also called Meltemi, blow between NW and NE in the Aegean Sea in the summer season. This is a consequence of the low-pressure system that stretches from Anatolia to north-west India that is formed by intense heating of the region (a "heat-low"). This peak in August when the mean wind speed is around $8\,\mathrm{ms}^{-1}$ and the mean of the maximum values is included between $15\,\mathrm{ms}^{-1}$ and $17\,\mathrm{ms}^{-1}$ in the southern Aegean. Although they produce choppy conditions for much of the time, their principal benefit is to moderate temperatures in coastal regions. Occasionally, they may produce sudden wind and squall however be associated with violent thunderstorm; this is the most dangerous situation for the Aegean Sea and Saroniko Gulf and occur when the circulation pattern combines with an influx of cold air at upper levels from more northerly latitudes while at the surface the hot air continues to be drawn round the heat low. This combination makes the atmosphere highly unstable and hence ideal

for convective activity and the formation of vigorous thunderstorms (Burroughs and Lynagh, 1999).

Besides the Saroniko Gulf that is between Attica peninsula, Corinth and Peloponnesian regions as illustrated in Fig. 1, have in the summer period a most complicated atmospheric circulation. This is caused by the formation of the sea breeze that affect the west coast of the Attica peninsula from the direction of southwest. These two winds, the Etesian and the sea breeze that blow in opposite directions, can converge in the coastal zone of the Attica peninsula thus the speed of the gradient wind can fluctuate in function of the speed of the thermal wind.

The purpose of this work is to study the Etesian winds in the Saroniko Gulf to understand the complicate physical processes that generate the windstorms in this zone. Therefore it is developed a detailed study of the weather patterns that characterize the Etesian winds and the sea breeze in the Saroniko Gulf and a comparisons of the wind data measured on land by the station of Helliniko airport (years: 1990–2000) and at sea by a buoy (years: 2000–2002) during the month of August. It is presented a statistical model that shows, as a probability of occurrence, the variations of the wind speed in function of the variations of the wind direction. This model can be used to make a now casting of the Etesian winds and to detect the chance to have windstorms in the Saroniko Gulf. Finally the statistical model is compared with the RAMS limited area model (Pielke, 2002) during a test case in the August 2004 (16/08/2004); the obtained results of the now-casting using these two different methods are compared to data measured by station of the automatic weather network of the HNMS.

2 The weather patterns in the Saroniko Gulf

Following the studies of Mayencon (2004) it should be expected to face up with two weather patterns: the Etesian pattern (Meltemi wind) and the sea breeze pattern.

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Fig. 1. Saroniko Gulf. The flags show the principal direction of the sea breeze from SW and of the Etesian wind from NE (courtesy: Microsoft Encarta).

2.1 The Etesian pattern

Between a trough located on the south and south west of Turkey, and high pressures on the north of the Aegean sea and on the Balkanize peninsula, a thermal flow can establish from N to NE. The thermal depression which goes through the Peloponnesian in the afternoon does not affect the Saroniko Gulf if the N–NE wind in upper atmosphere is higher than $6 \div 7 \, \mathrm{ms}^{-1}$. The sounding of Hellenikon Airport, very closed to the site, show that, in the Etesian pattern, the wind in upper atmosphere at 800 m is higher than $7 \, \mathrm{ms}^{-1}$ and always blows from N or NE. This is a typical characteristic of an Etesian weather.

The Etesian wind blowing from N–NNE in the Aegean sea, often strong, turns NE and blows slowly in the region of Athens. In the surface, in the morning, the wind is generally weak and blows N–NW. It becomes stronger, veering, reaching NNE or NE in sea, even ENE. Going down the hills, it is an irregular wind in intensity and in direction. In the late afternoon, this wind usually backing to north. On the coast closed to Athens, the Etesian who blows from land is often warm in summer, around 35°C in the afternoon, reaching sometimes 40–43°C, while the sea temperature is around 25°C. This temperature gap creates just above the sea a thin air layer in which the temperature grows with the altitude

(thermal inversion). This thin layer very stable slows down strongly the Etesian on sea, even very closed to the coast. This phenomenon of thermal inversion explains the generation of areas of calm on sea.

When the Etesian is weak (less than $5 \, \mathrm{ms}^{-1}$ in upper atmosphere) and as it is warm, it cannot prevent that the sea breeze blow from SW and that the wavy breeze front reach the south-west coastal zone of Attica. This zone of convergence, where the Etesian wind and the sea breeze meet, may reach $1 \div 2$ km. In the late afternoon, the decrease of the temperature leads to a strengthening of the Etesian, which pushes back the sea breeze front in the open sea.

2.2 The sea breeze pattern

When the gradient wind is absent or weak and when the onshore convection is sufficient, the wind in the morning blows feeble from NW or WNW, then it strength and back to W or WSW around 12:00 LT or 13:00 LT. It goes on progressively backing until 15:00 LT, to reach SW or SSW. Between 15:00 LT and 17:00 LT, when the pressure in Athens-Hellenikon becomes lower of 2 hPa than the pressure at the Sounion Cape, generally the wind turns quickly to the SE. This rotation is due to the important landscape of the peninsula at the SE of Athens, which canalizes the marine breeze. When the breeze generated by the thermal trough centered

Fig. 2. Position of the Helliniko Airport AWS and the buoys in the Saroniko Gulf.

on the Peloponnesian arrives, the wind turns S and become stronger.

2.3 Forecast of the weather pattern in the Saroniko Gulf

The identification of the precursor and large-scale synoptic conditions in the Mediterranean sea have been the subject of specific interest in the last few years to analyze the extreme precipitations events (for a review of the studies see Rudari et al., 2002). In our case it is very important to detect the weather patterns to forecast the wind conditions in surface and to detect the position of the convergence zone between the Etesian wind and the sea breeze. To define the weather patterns in the Saroniko Gulf two precursors used. The first was the wind speed at 925 hPa (approximate altitude of 800 m). In effect, as shown by Mayençon (2004):

- $u_{925} \leq 6.5$ ms^{-1} from NNE-NE: the sea breeze was the wind in the Saroniko Gulf near the coastal zone of Attica.

- 6.5 ms$^{-1} < u_{925} < 8.5$ ms^{-1} from NNE-NE: in the Saroniko Gulf near the coastal zone of Attica was the Sea Breeze Convergence Zone (SBCZ).

- $u_{925} \geq 8.5$ ms^{-1} from NNE-NE: the Etesian wind was in the Saroniko Gulf.

The second precursor of the weather pattern was the "severe weather index" k (George, 1960)

$$k = (T_{850} - T_{500}) + T_{d850} - (T_{700} - T_{d700}) \qquad (1)$$

in which T is the air temperature °C and T_d is dew point °C; these temperatures are referring to particular pressure levels (850 hPa, 700 hPa and 500 hPa). Usually the k index is used to understand the stability of atmosphere (Hand et al., 2004) and the $k \geq 20$ is a signal of risk of thunderstorm. In the Saroniko Gulf is possible to use the k index to define the characteristic weather patterns of the race area. It is true that in August, in the Saroniko Gulf, the air is dry and so, also if the k index is higher (risk of thunderstorm), it is not possible to have clouds. For this reason it is suggested to use the k index, in the Saroniko Gulf, to define the weather patterns and not to define the atmospheric stability. In the forecast, issued during the Olympic Games (August, 2004), was used the k index to define the weather patterns and the "relative vorticity" [10^{-5} s^{-1}] (AA.VV., 2003) to decide the state of atmosphere.

In the Saroniko Gulf, were identified the following weather patterns:

- $k \leq 15$°C: "light" sea breeze (wind speed 2–3.1 ms^{-1}) or Etesian wind (see the wind speed in upper atmosphere at 925 hPa).

- 15°C$ < k \leq 20$°C: "medium" sea breeze (wind speed 3.1–4.2 ms^{-1}).

- 20°C$ < k \leq 27$°C: "medium" sea breeze (wind speed 3.1–4.2 ms^{-1}) and possible off-shore gradient wind (from SW).

– 27°C<k≤33°C: Sea Breeze Convergence Zone (SBCZ).

– k>33°C: "strong" sea breeze (wind speed 4.2–6.7 ms^{-1}).

During the forecast issued in August 2004 for the Olympic Games the skill of this parameterization of the weather patterns was around 75%.

3 The now-casting in the Saroniko Gulf

In this section the methods to build the statistical model, also called "swing table" (Molyneux, 2003), used to prepare the forecast during the August 2004 will be analyzed. In other the same model, merged with a meso-scale non-hydrostatic model (RAMS), can used to issue the now-casting warning for the situation with high speed of the wind. The last part of this section will be dedicated to study the test case of the 16/08/2004 in the Saroniko Gulf.

3.1 The database

The first step to build the statistical model ("swing table") is to research the reliable meteorological data measured by a significant weather station. The goal of this research is to make the statistical model for the August month. The Hellenic Meteorological Service (HNMS) in cooperation with the Hellenic Center for Marine Research have managed many Automatic Weather Stations (AWS) and three buoys located in the Attica Peninsula and in the Saroniko Gulf near the coastal zone (http://www.hnms.gr/hnms/english/index_html). The analyzed database is represented by the meteorological data measured from the AWS of the Hellinikon Airport and the buoy deployed at 37.85 N–23.70 E (see Fig. 2). The buoy has measured in the months of August of the years 2000, 2001 and 2002 (http://www.sailing.org/menu. asp?MenuID=tev'GBv,zBpOvIJ9SuH8NX6/T'b3EVKHi, XNEnGMSgr7UTWdXfjhTg5Rtj10cizf) but the data are recorded only every three hours (average hourly of the speed and the direction of the wind), so the database of the buoy was absolutely not significant. Nevertheless the Hellinikon Airport AWS have recorded the weather data in the August's months between the 1990 and the 2000 every hours (average hourly of the speed and the direction of the wind).

It is evident that the crucial year is the 2000 as there are the contemporary measures of the buoy and the Hellinikon Airport AWS. It is possible to transpose the onshore database to the sea if the equations that link the onshore wind speed and direction to the offshore wind speed and direction are known with an acceptable correlation coefficient. By this way, it can work on a virtual database for the Olympic Games period and achieve to release the swing tables.

3.2 Transposition of the wind speed

The first effort was to research the equations that correlate the land-based wind data with the off-shore database. Therefore, using the theory of transposition (Hsu, 1988) and fixed the condition r^2>0.6, it was obtained for the wind speed (Pezzoli and Dublanc, 2004b):

$$U_{sea} = 0.225 + 0.771 U_{land} \qquad (2)$$

where the U_{sea} is the wind speed in the sea at buoy position and the U_{land} is the wind speed in the land at Hellinikon Airport AWS. Obviously, this equation is valid as well in m/s as in knots (kts) and the correlation's coefficient was r^2=0.652.

3.3 Transposition of the wind direction

To transpose the wind direction the first objective was to generate a model valid in each wind conditions. It was decided to divide the problem in two parts, depending on the direction of the wind recorded in the Hellenikon Airport AWS. Two great weather patterns are clearly identified:

– the Etesian pattern, with a wind blowing from North.

– The sea breeze pattern, with a wind blowing from South.

These typical situations are explained in detail in the Sects. 2.1 and 2.2. For the Etesian pattern the wind recorded at Hellenikon Airport AWS was between 320° and 140°; on the contrary for a wind between 140° and 320°, the see breeze pattern was prevalent.

Besides it was founded a relation between the wind direction onshore and offshore. It can assume that the difference between onshore and offshore direction counted in the clockwise angle is function of the temperature gap between air and water (Hsu, 1988). The form of this equation was:

$$\Delta\theta = \zeta + \eta\Delta T \qquad (3)$$

where ζ, η are constant values to determine, ΔT is the air-water temperature difference (°C) and $\Delta\theta$ is the clockwise angle between over-land and over-sea winds (degrees). However, when the temperature gap is above 5°C, it is bound to bring some corrections by including a dependence of the offshore wind speed:

$$\Delta\theta/\Delta T = \chi(U_{sea}) \qquad (4)$$

where χ is a function, which must be found empirically.

3.3.1 Correlation of wind directions for the Etesian pattern

Also in this case r^2 is fixed as r^2>0.6 to have an acceptable model to transpose the wind direction from the land to the sea. Analyzing the meteorological database, it was subdivided the case when the ΔT <5°C and the ΔT >5°C. For a temperature gap under 5°C, the model follows the theory

Table 1. Example of "swing table".

Table 2. Forecast of the average wind direction and speed for the 16/08/2004 and measures of the buoys A, B and C in the Saroniko Gulf.

hour	forecast (RAMS and "swing table")	Buoy A	Buoy B	Buoy C
[LT]	[°/ms-1]	[°/ms-1] 37.9 N–23.65 W	[°/ms-1] 37.88 N–23.7 W	[°/ms-1] 37.82 N–23.7 W
09:00	350°/4.633	014°/6.528	006°/8.145	356°/7.734
10:00	355°/5.148	350°/5.156	019°/7.734	008°/8.086
11:00	000°/5.662	021°/8.034	028°/9.785	016°/8.906
12:00	000°/6.178	023°/9.258	022°/12.129	020°/10.020
13:00	010°/7.722	026°/8.437	016°/10.020	011°/8.906
14:00	015°/9.781	015°/6.898	018°/9.375	008°/8.262
15:00	017°/10.296	008°/9.855	020°/11.367	004°/9.844
16:00	016°/9.781	016°/9.699	013°/8.906	013°/9.961
17:00	015°/8.236	025°/11.459	018°/9.961	010°/8.730
18:00	005°/6.178	015°/8.005	005°/9.668	001°/8.730

(see Eq. 3) as it is obtained a linear relation with a coefficient of correlation r^2=0.642, in the range of acceptability:

$$\Delta\theta = -72.585 + 16.662\Delta T \qquad (5)$$

For a $\Delta T > 5°C$, in order to have a more reliable model, it is subdividing the ranges of temperature and it is searched a first equation for $5 < \Delta T < 8°C$ and a second equation for $\Delta T > 8°C$, because it is noticed that in the Etesian pattern ΔT may reach $12 \div 13°C$. Considering this subdivision, it is arrived a following results:

– $5 < \Delta T < 8°C$:

$$\Delta\theta/\Delta T = -2.187U_{sea}^2 + 20.931U_{sea} - 47.146 \qquad (6)$$

with r^2=0.957.

– $\Delta T > 8°C$:

$$\Delta\theta/\Delta T = -0.0295U_{sea}^4 + 0.746U_{sea}^3 - 7.140U_{sea}^2$$
$$+28.701U_{sea} - 42.023 \qquad (7)$$

with r^2=0.684, always in the range of acceptability.

3.3.2 Correlation of the wind directions for the sea breeze pattern

As it is said in the Sect. 3.3, for a $\Delta T < 5°C$, it is searched for a linear model. It is obtained:

$$\Delta\theta = -53.37 + 9.962\Delta T \qquad (8)$$

with an r^2=0.945 which is a great result.

For $\Delta T > 5°C$, it should look for a velocity correction. The first results obtained were bad and it should look for another
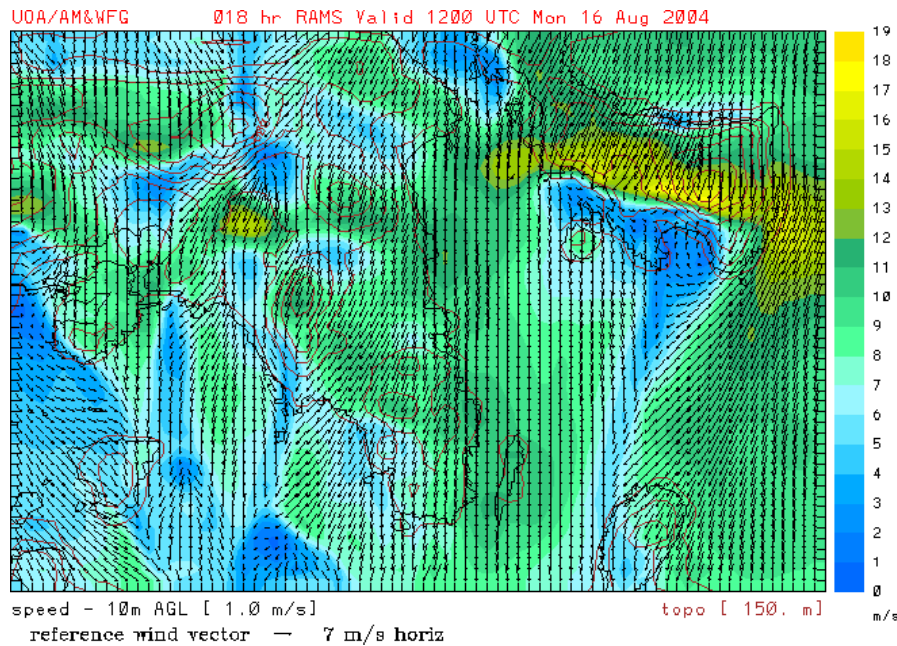
Fig. 3. Forecast of the surface wind from RAMS model for the 12:00 UTC (15:00 LT) of the 16/08/2004 (courtesy: University of Athens).

empirical model since the model with velocity correction was not valid. At this point of the study, it was forced to pass upon the theory and try to find out a model that respects the best the values recorded by the Hellinikon Airport AWS and the buoy. It was managed to find out a dependence of the square of the gap of temperature, so the model for this case was:

$$\Delta\theta = -19.766 + 73.328\Delta T - 8.422\Delta T^2 \qquad (9)$$

with an amazing r^2=0.999.

After to have transposed in the sea the database of Hellinikon Airport AWS using the equations determined in the Sects. 3.3.1 and 3.3.2, it was checked the results obtained in order to erase those who are unrealistic. To have a better overall view, it was bound to create a table for every year (1990–2000) in which was reported the unrealistic situations. These tables have in input the days, the hours of the August months and the transposed meteorological data; the aim of these tables is to provide an average rate of reliability, that is to say with which percentage the model of transposition is valid. In according to the results found in these reliability tables, the model of transposition, represented by the Eqs. (5)–(9), gives "realistic" results with an average rate of 80%. It may assume that the equations established in the Sects. 3.3.1 and 3.3.2 is reliable, considering the complexity of the meteorological phenomena in the Saroniko Gulf, so it is possible to affirm that also the new database build for the sea is valid. It is important to remember that the transposition's equations obtained are valid for the August months.

3.4 The statistical model (the "swing table")

The first step to build the "swing table" is to collect all the reliable data, from 1990 to 2000 during the August's month,

of the virtual buoy (wind speed and direction) at h and $h+1$. Then a MATLAB program calculates all the changes in wind direction and speed between h to $h+1$. The result is a matrix in which appears the variation of the wind speed in the first column and the change of the wind direction in the second column. The number of the rows is equivalent to examined number of data.

Afterwards the relative frequency is calculated as:

$$f_{r_c} = \frac{n}{N} \qquad (10)$$

where n is the number of the cases of the variations in each considered class of wind speed and wind direction and N is the number of the examined data coincident with the rows of the above-mentioned matrix. The final results is the "swing table" that is a matrix 13 by 3, which represents 39 conditions of wind speed and direction (see the Table 1 for an example). Three conditions for the wind speed are considered:

– increase: from h to $h+1$, wind speed has won more than 1 knot.

– Steady: changes from h to $h+1$ remain between −1 knot and 1 knot.

– Decrease: from h to $h+1$, wind speed has lost more than 1 knot.

For the direction, (calculated on the Geographic North), changes of 10° are considered. Changes above 50° and under −50° are in the rows shift right and shift left in grey. The numbers in bold and italic in the cells are the percentage of occurrence of one of the 39 wind conditions (equivalent to $f_{r_c}\times100$). In red is the percentage of backing, between −50°
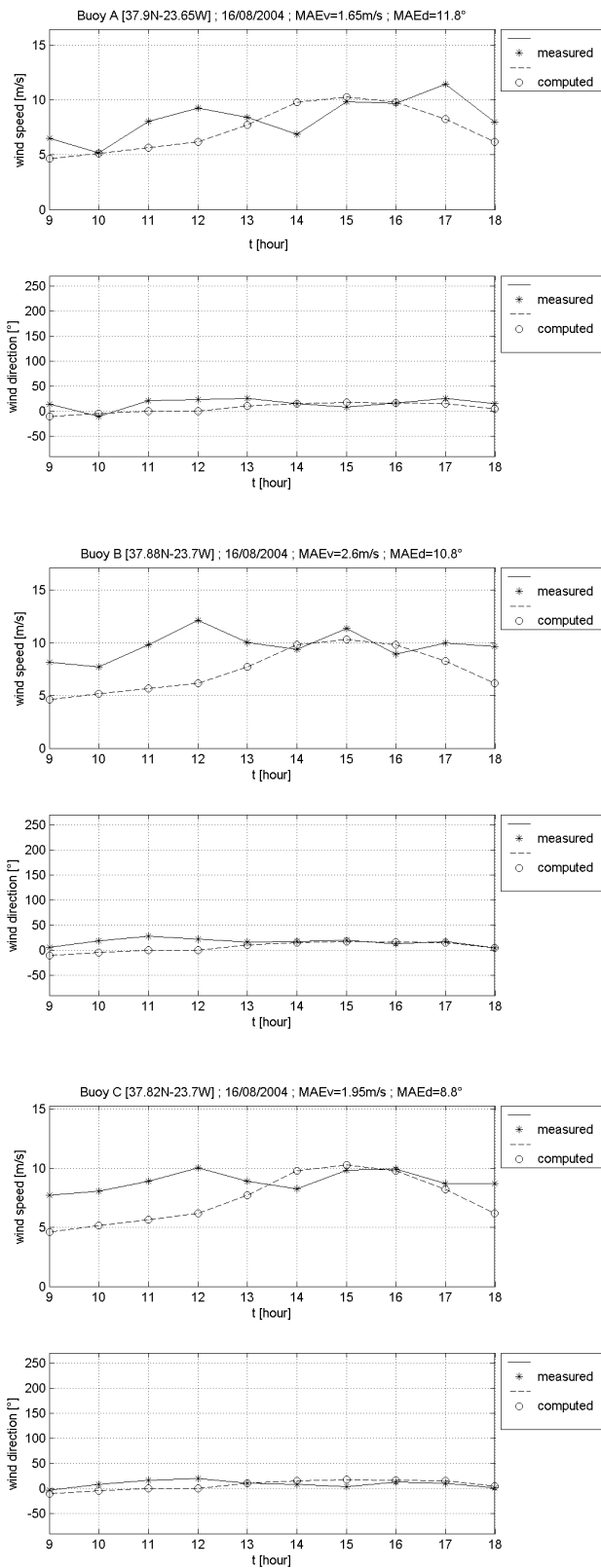
Fig. 4. Comparison between forecasted (computed) and measured wind in the buoys.

and −10°, and in green is the percentage of veering, between 10° and 50°; in yellow and orange, the situations which happen the most frequently. At last, the approximations have been made as follows: a change of 14.9° is counted like a change of 10°. A change of 15.1° is counted like a change of 20°. The 15° remain in the row change of 10°.

This new concept of statistical model will help the forecasters to issue the wind-warnings in the now-casting and short-term mode. Using the real-time measures of the shift of the wind direction, it will be possible with the "swing table" (see Table 1) to detect the probability to have of the variations of the wind speed.

3.5 The case studies: 16 August 2004

The 16 August 2004 the high pressure was positioned between the Black Sea and the west zone of the Mediterranean Sea and a low pressure was located in the Turkey (http://www.wetterzentrale.de/topkarten/fsfaxsem.html). In the meso-scale analysis was possible to observe an high pressure in north of Aegean Sea and low pressure in the south of Peloponnesian; this situation generated a strong flux from N in the Saroniko Gulf. In the same time with the passage in the morning of the instability line in Turkey and in the west of Peloponnesian were associated more Cumulus and Cumulonimbus clouds; the forecasted k index was equal to 30.8°C and represented a convergence zone, this forecast was valid for the morning but not after 11:00 LT when the Etesian wind was fully developed in the Saroniko Gulf (see Table 2). The forecasted wind in upper atmosphere (925 hPa) was from 22° with a wind speed of $14.4\,\mathrm{ms^{-1}}$ that confirm the Mayençon's theory (see Sect. 2.3). Also the RAMS model (http://forecast.uoa.gr/) showed a day of Etesian wind with a forecast of the increasing of the wind speed between 13:00 LT and 14:00 LT (see Fig. 3); in effect the buoys have measured the strengthening of the wind after the 11:00 LT (see Table 2). Nevertheless the "swing table" model simulated well the increasing of the wind between 12:00 LT and 14:00 LT when, for an oscillating wind or for a veering, the probability to have an increasing of the wind was altogether equal to 24% (Pezzoli and Dublanc, 2004a). For the time interval included between 14:00 LT and 17:00 LT the same model showed us a good results: in effect the more higher values of the probability was for a steady wind in the direction and in the speed. Finally between 17:00 LT and 18:00 LT the graphics, represented in the Fig. 4, showed us a backing and a decreasing of the wind; in this case the statistical model provided an 8% of probability to have the decreasing of the wind's speed associated a the left shift as against the right shift that had the 5% of probability.

The good reliability of the forecast is demonstrated by the Mean Absolute Error, also called MAE (Wilks, 1995), which is included between $1.65\,\mathrm{ms^{-1}}$ and $2.6\,\mathrm{ms^{-1}}$ for the wind speed and between 8.8° and 11.8° for the wind direction. The maximum MAE of the wind speed is in the Buoy B, but this error is lower than 30% in comparison with the average of the measured wind speed.

4 Conclusion

In conclusion it can affirm that the methodology proposed in this research to issue the forecast is reliable. It is important to consider that all the three methods (weather patterns, "swing table" with the probability of occurrence and atmospheric models) have to be used contemporary to prepare a forecast with the best skill. Finally it can affirm that these three methods are statistical (weather patterns and "swing-table") and deterministic (atmospheric models) and they merge in the unique model called "mixed statistical-deterministic". Therefore the forecast is written with the "Reliability Index" that shows the probability of occurrence of the same forecast; this index is used as decision-maker to build the economic estimation of the weather forecast (Pezzoli and Franza, 2003).

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