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FOCA: a new quality-controlled database of floods and catchment descriptors in Italy

Pierluigi Claps, Giulia Evangelista, Daniele Ganora, Paola Mazzoglio, and Irene Monforte

Department of Environment, Land and Infrastructure Engineering, Politecnico di Torino, Torino, 10129, Italy

Correspondence: Pierluigi Claps (pierluigi.claps@polito.it)

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Abstract. Here we present FOCA (Italian FLOod and Catchment Atlas), the first systematic collection of data on Italian river catchments for which historical discharge time series are available. Hydrometric information, including the annual maximum peak discharge and average daily annual maximum discharge, is complemented by several geomorphological, climatological, extreme rainfall, land-cover and soil-related catchment attributes. All hydrological information derives from the most recently released datasets of discharge and rainfall measurements. To enhance the reproducibility and transferability of the analysis, this paper provides a description of all the raw data and the algorithms used to build the basin attribute dataset. We also describe the approaches adopted to solve problems encountered during the digital elevation model elaboration in areas characterized by a complex morphology. Details about the data quality-control procedure developed to detect and correct errors are also reported. One of the main novelties of FOCA with respect to other national-scale datasets is the inclusion of a rich set of geomorphological attributes and extreme rainfall features for a large set of basins covering a wide range of elevations and areas. Using this first nationwide data collection (available at <https://doi.org/10.5281/zenodo.10446258>, Claps et al., 2023), a wide range of environmental applications, with a particular focus on flood studies, can be undertaken within the Italian territory.

1 Introduction

The availability of complete, updated and quality-controlled hydro-geomorphological information for areas nationwide is a key requirement for a vast range of applications from hydrological modeling to hydraulic simulations and rainfall-runoff analyses. Hydro-geomorphological catchment information, also called catchment attributes, can also provide a comprehensive description of the landscape and of how the catchment stores and transfers water (Addor et al., 2017).

The work of Linke et al. (2019), who created the HydroATLAS database by mapping key hydrological variables at 500 m grid resolution, represents a considerable effort to provide hydrological, climatological and land-cover information on a global scale. In that work, however, few hydrological and climatic attributes were considered, and geomorphological catchment descriptors are completely missing from it. In Europe, the importance of river basin catalogs acting as a support to water resources management

was strongly stressed by the publication of the Water Framework Directive in 2000. River catchments that cross administrative boundaries can suffer from a lack of incomplete and non-uniform information. To tackle this shortcoming, in early 2000, an initiative undertook the creation of a pan-European river and catchment database (Vogt et al., 2007). In recent years, considerable efforts have been dedicated to producing similar datasets at the national scale for several European countries. Relevant examples are the datasets built in the framework of the CAMELS (Catchment Attributes and MEteorology for Large-sample Studies) initiative, as in the cases of Great Britain (Coxon et al., 2020), France (Andréassian et al., 2021; Delaigue et al., 2022), Switzerland (Höge et al., 2023) and Germany (Loritz et al., 2022). Similar datasets from other continents that were developed within the same framework have also been released, such as those for the United States (Addor et al., 2017), Chile (Alvarez-Garreton et al., 2018), Brazil (Chagas et al., 2020)

and Australia (Fowler et al., 2021). Most of those datasets have since been linked together in order to build the Caravan dataset (Kratzert et al., 2023). Previous notable examples are the atlases and datasets developed in Switzerland (<https://hydrologicalatlas.ch/>, last access: 11 March 2024), Austria (Fürst et al., 2009) and Canada (Arsenault et al., 2016). In recent years, other large-scale datasets have been developed independently of the CAMELS framework, such as those related to China (Hao et al., 2021) and to the area of the upper Danube up to the Austria–Slovakia border and some nearby catchments (Klingler et al., 2021).

As of today, only partial-coverage datasets (in terms of both spatial extent and the number of variables) are available for Italy. We can mention the datasets developed for the northwest of Italy (see, e.g., Barbero et al., 2012; Gallo et al., 2013), the northeast of Italy (Crespi et al., 2021) and North-Central Italy (Pavan et al., 2019) as well as some studies that aimed at mapping a few variables across the entirety of Italy (ISPRA, 2005a, b; Claps et al., 2008; Crespi et al., 2018; Braca et al., 2021). Among the above examples, the database realized within the CUBIST (Characterisation of Ungauged Basins by Integrated use of hydrological Techniques) project (Claps et al., 2008) was the only attempt to build a multi-variable national data collection; all the other collections focus on only one specific topic (rainfall or discharge or geomorphology). A complete and updated database of hydro-geomorphological variables relating to the main catchments of Italy is therefore still missing. One of the reasons behind this gap is the dismantlement of the National Mareographic and Hydrographic Service (Servizio Mareografico e Idrografico Nazionale, SIMN), which has led to the federated management of the national monitoring network by 21 different administrative agencies. Nonetheless, the high network density of the rain and stream gauges that are available throughout the country allows the compilation of nationwide catalogs, which can also take advantage of the considerable number of local studies performed using systematic samples of hydrological measurements.

A recent effort to collect relevant hydrological information with national coverage is represented by the work of Claps et al. (2020a, b, c), in which annual maximum peak discharges (hereafter named “peak discharges” for brevity) and average daily annual maximum discharges (hereafter named “daily discharges” for brevity) from 1911 to 2016 for 631 Italian catchments are published, together with some basic geomorphological catchment attributes. This collection was mainly created to integrate previously untracked flood discharge measurements whose historical records were not available in hydrological yearbooks; they were only available in special publications. In this paper, we aim to substantially enhance the existing catchment and peak discharge catalogs by complementing the hydrological data with a large set of climatological, geomorphological, soil and land-use attributes, many of them computed from recently released databases. All data sources used comply with the follow-

ing criteria: (a) nationwide coverage; (b) consistency in data quality (i.e., no regional or local biases); and (c) adequate original resolution in relation to the type of information. All the data sources used in this work can be downloaded from public repositories, except for a few variables that are not easily accessible. In these cases, the original data sources were included in the dataset created in this work, together with the mean catchment descriptors, to allow replicability. The result of this work is FOCA (Italian FIOod and Catchment Atlas), the first national-scale catchment attribute dataset in Italy (Claps et al., 2023).

The paper is structured as follows. In Sect. 2, the database history and the rationale used to select catchments to include in the database are described. In Sects. 3 and 4, we present the different categories of raw data (which were already partly quality-controlled and only required some adjustments), which largely relate to digital elevation models and streamflow measures, and partly processed data, which derive from the aggregation and manipulation of rainfall and soil information. In Sects. 3 and 4, we also provide an overview of all the geomorphoclimatic attributes computed for each catchment (together with the methodologies used for their definition and the algorithms used for their evaluation) to grant replicability and thus allow researchers to perform the same study or extend this one to other catchments. In Sect. 5 the data availability is presented. In Sect. 6, the main characteristics of the FOCA dataset are summarized, and some concluding remarks are presented.

2 Presentation of historical catalogs and the rationale for catchment selection

Since the early 1990s, the Italian monitoring network was managed by the SIMN. This institution was in charge of the collection and validation of the data and its publication in a series of hydrological yearbooks with yearly updates. Two standardized documents (part I and part II) were usually published each year. The first volume (part I) was dedicated to temperature and precipitation; the second volume contained information and elaborations related to daily and monthly average discharges.

Due to this separation, to retrieve an entire time series measured by a gauging station, it was necessary to consult all the yearbooks. All the mentioned yearbooks are available as images (in gif format) at <http://www.bio.isprambiente.it/annalipdf/> (last access: 23 October 2023). To facilitate access to the discharge data, this information was then processed by the SIMN at monthly time resolution and grouped by individual gauging station with the aim of providing several years of data on one data sheet for each station. This summarized information, available up to 1970 and complemented with some basic catchment information, became *Pubblicazione no. 17 (Publication no. 17)*; this also included the annual peak values that were not directly available in the yearbooks.

This document was updated about every 10 years from 1934 on (several issues are still available) by progressively including the data acquired up to 1970 (Servizio Idrografico, 1980).

Unfortunately, the updating of this publication was interrupted, as the SIMN was dismantled about 30 years ago and 21 different local agencies then became responsible for the management of the hydrographic services. This interruption of the publication of *Pubblicazione no. 17* had a particularly negative effect on access to the peak discharge data. They were only published in *Pubblicazione no. 17* as the maxima of the instantaneous discharge and have never been available in the hydrological yearbooks. This means that, for many years, peak discharges measured after 1970 were only available in the regional hydrographic offices. The main effort to recover these unpublished data was made by the GNDICI (Gruppo Nazionale per la Difesa dalle Catastrofi Idrogeologiche) as part of the VAPI (VALutazione delle Piene) project (<http://www.idrologia.polito.it/gndci/Vapi.htm>, last access: 23 October 2023), which was subsequently integrated into the CUBIST project.

The first major advancement in the creation of a nationwide database was achieved by the CUBIST project (Claps et al., 2008). More specifically, historical information previously available only in printed format was digitized and merged with more recent data. Moreover, each catchment closed by the gauging stations included in this collection was characterized from a climatological and geomorphological point of view by means of the computation of key catchment attributes.

In the following years, some regional-scale works were conducted to update this systematic collection but without nationwide coordination. For example, in Northwest Italy, the first follow-up work was the *Catalogo delle portate massime annuali al colmo del bacino occidentale del Po* (Barbero et al., 2012), which includes a considerable number of additional observations from the gauging stations operated by the regional environmental agency as well as additional gauging stations managed by other public or private bodies, such as Enel (a major Italian energy provider) and the CNR (Italian Research Council). In this publication, a systematic and extensive data validation was performed to convert into discharge values the hydrometric stage values that had been recorded by the SIMN but had never been processed with adequate rating curves. This collection summarizes the information for 140 catchments with at least 5 years of data on peak or daily discharges, and it contains some basic information on the catchments. Unfortunately, this publication does not contain all the attributes evaluated within the CUBIST project. This regional collection was thus improved with the release of the *Atlante dei bacini imbriferi piemontesi* (Gallo et al., 2013), which contains geomorphological attributes of about 200 gauged river catchments located in the same region. In this case, however, the peak discharges were not reported, as they were not available for all the basins. Similar works were also carried out for other limited areas in the

framework of local studies on flood frequency analysis (e.g., Rossi and Caporali, 2010).

Such collections of flood datasets were recently homogenized, integrated and updated by including the most recent data acquired across Italy, leading to the release of a comprehensive catalog of floods (*Catalogo delle Piene dei Corsi d'acqua Italiani*; Claps et al., 2020a, b, c). However, in this case, only basic geomorphological information was included for the considered basins.

The catchment selection criterion adopted in the present work stems from the purpose of improving upon the work reported in Claps et al. (2020a, b, c). The 631 chosen catchments are those for which peak or daily discharges are available and are therefore all included in Claps et al. (2020a, b, c). Details on the different sources of the historical discharge time series used and their integration are given in Sect. 4.4.

3 Data and catchment geomorphological attributes

3.1 The digital elevation model

As of today, the only national-scale high-resolution digital elevation model (DEM) available for the whole of Italy is TINITALY/01 (Tarquini et al., 2007). Despite its very high spatial resolution (10 m) – which makes its use quite interesting, especially for the delineation of small catchments – it presents the drawback of being obtained by merging separate DEMs of single administrative regions, so it does not allow us to work with the same accuracy level nationwide. To overcome this drawback, we adopted the DEM from the Shuttle Radar Topography Mission (SRTM) at 30 m spatial resolution (Farr et al., 2007).

The processing of the DEM and, in general, all the other catchment attributes was performed with open-source software, namely GRASS GIS and R. In this work, the original SRTM DEM was re-projected into the WGS84/UTM zone 32N coordinate system by means of a bicubic interpolation and resampled to obtain integer cells, i.e., the data format requested by the GRASS GIS add-on *r.basin* (Di Leo and Di Stefano, 2013) that was used to derive most of the geomorphological attributes, as will be discussed in the following subsections. DEMs usually contains pits (i.e., elevation values that are much lower than those of nearby pixels; these are errors due to the resolution of the data) that should be filled to ensure the proper delineation of catchment boundaries and drainage networks. Thus, the pit-filling procedure was carried out using the *r.hydrodem* add-on for the GRASS GIS (Lindsay and Creed, 2005).

The *r.basin* add-on algorithm requires a series of information before it can be executed. The input parameters required by *r.basin* for the attribute extraction routine are presented in the following subsections.

3.2 Catchment boundaries

To determine the basin boundaries, it is first necessary to generate the drainage network from the depitted DEM. This step can be done only once with the *r.basin* command and involves the following steps: (i) the calculation of the drainage directions with the multiple flow direction (MFD) algorithm; (ii) the calculation of the flow accumulation, i.e., the total contributing area (TCA) map; and (iii) the estimation of the stream network after specifying a threshold value that defines the minimum drainage area required to initiate a channel. In this work, threshold values equal to 0.02, 0.1, and 1 km² are used to extract the stream network for basin areas smaller than 1 km², basins between 1 and 10 km², and basins larger than 10 km², respectively. These values were identified considering recommendations derived from several research works that investigated the spatial resolution sensitivity of catchment geomorphologic properties and its effect on the hydrological simulation, such as Montgomery and Foufoula-Georgiou (1993), Yang et al. (2001), and Beighley and Gumadi (2011). Those works highlighted that river networks generated with larger threshold areas tend to lose relevant information, which led us to avoid using a fixed threshold value, as it could be problematic for small basins.

To proceed toward delineation, a unified determination of the basin outlet coordinates is necessary, as the official coordinates taken from hydrometric stations do not necessarily coincide with stream locations automatically determined from the DEM. So, the coexistence of two sets of coordinates, the real and the operational DEM-based ones, must be properly accounted for during the creation and management of the dataset. The second set of coordinates, the DEM-based ones, are evaluated on the basis of the reprojected SRTM DEM; were a new DEM to be used in future works, this set of points would need to be evaluated again. Even though the *r.basin* algorithm is able to automatically work with input outlet coordinates that do not exactly overlap with the DEM-based river network by snapping to the closest point belonging to the network, in some cases the relocation may fail and manual repositioning of the outlets will thus be required. Compared to the real coordinates, the required adjustments are, in some cases, of the order of a few kilometers. This adjustment is needed to obtain a river network that matches with the reference one provided by the Istituto Superiore per la Ricerca e la Protezione Ambientale (ISPRA, available at <http://www.sinanet.isprambiente.it/it/sia-ispra/download-mais/reticolo-idrografico/view>, last access: 23 October 2023).

As expected, the procedure for catchment boundary delineation is generally very accurate where the differences in elevation are quite marked (i.e., in alpine areas). Thus, the only manipulation required is the repositioning of the catchment outlet, which, in our case, is always positioned at the gauging stations. However, this method does not always provide the real drainage directions in flat areas, and it is there-

fore necessary to manually force the DEM to correct the total contributing area (TCA) map built by the *r.basin* procedures, a practice commonly known as *stream burning* (Lindsay, 2016). This operation is performed individually, when needed, by comparing the unconstrained river network produced by *r.basin* with the reference one provided by ISPRA. This quality checking was iteratively repeated by carving rivers or inserting artificial barriers to force the stream onto the correct path. The extent of this operation, i.e., the location(s) and length(s) of the carved or (and) walled portion(s), is unique for each area; for this reason, this operation requires an individual assessment. The procedure described above is depicted in the flow chart in Fig. 1.

The catchment boundaries resulting from the delimitation are finally made available in vector format in the WGS84 UTM32 N (EPSG 32632) coordinate system. The catchment boundaries were used as masks to clip several layers of climatological and soil-related attributes. Moreover, the provision of this geographic information will allow users to expand this database by computing other descriptors using possible new gridded datasets.

3.3 Geomorphological catchment attributes

Geomorphoclimatic information, along with soil characteristics, is essential when trying to characterize how the catchment stores and transports water both on and below the surface. In this work, we first selected from the literature a large set of morphological attributes that can be directly obtained by processing the 30 m resolution STRM DEM. Automatic procedures based on the GRASS add-ons *r.basin*, *r.stat*, *r.slope.aspect*, *r.stream.stats* (Jasiewicz, 2021) and *r.accumulate* (Cho, 2020) are used for this computation. This leads to 61 geomorphological attributes (see Table 1 for a complete description).

The provision of this set of 61 different attributes for more than 600 catchments nationwide represents one of the strengths of the FOCA dataset. Comparing it to other relevant examples that we found in the literature, such as the CAMELS or the LamaH-CE datasets, one realizes that they include only a limited number of geomorphological descriptors and focus mainly on climatic, hydrologic, land-cover, soil and geological indices. As an example, in the CAMELS developed for the US (Addor et al., 2017) or for Brazil (Chagas et al., 2020), only basic topographic characteristics related to outlet coordinates, the catchment area, the mean elevation and mean catchment slopes are reported. In the LamaH-CE dataset (Klingler et al., 2021), some additional geomorphological attributes were included, such as the median basin elevation, the range of catchment elevations, the elongation ratio, the horizontal distance from the farthest point of the catchment to the corresponding gauge, and the drainage density. The inclusion of other descriptors in LamaH-CE was motivated by the need to know both the shape of the catchments and the stream network influence on

Table 1. List of geomorphological attributes, with a brief description of each and an indication of the algorithm/add-on used for its computation. All the attributes are computed by processing the SRTM DEM at 30 m resolution with the *r.basin* add-on, which takes advantage of other the GRASS GIS algorithms mentioned at the beginning of Sect. 3.3.

Attribute sub-category	Attribute	Notation	Units	Description
Altimetric and geometrical	Area	area	km ²	Catchment area computed by multiplying the area of a single pixel by the number of pixels within the catchment boundary.
	Mean elevation	elev_mean	m a.s.l.	Catchment mean elevation.
	Maximum elevation	elev_max	m a.s.l.	Catchment maximum elevation.
	Minimum elevation	elev_min	m a.s.l.	Catchment minimum elevation.
	Aspect	aspect	°	Mean of the angle of exposure in the horizontal plane of each cell of the catchment. The adopted convention is that north is 0° and the aspect is computed clockwise.
	Hypsographic curve	elev_x	m a.s.l.	Elevation values of the hypsographic curve (i.e., the curve that defines the distribution of catchment areas located within a specific elevation range). Each <i>x</i> corresponds to a different percentage of the area (2.5 %, 5 %, 10 %, 25 %, 50 %, 75 %, 90 %, 95 % and 97.5 %).
	Geographic centroid	x_g y_g	m	Coordinates of the pixel nearest to the centroid of the geometric figure resulting from the projection of the catchment on the horizontal plane.
	Length of the orientation vector	dir_length	km	Length of the segment linking the catchment centroid to the outlet.
	Orientation	orient	°	Angle of the orientation vector with respect to north.
	Mean slope 1	slope1	%	Mean slope value calculated by averaging the slope map.
Mean slope 2	slope2	%	Angle at the base of the right-angled triangle whose base is the square root of the catchment area and whose height is twice the median elevation of the catchment (relative to the closing section). This slope is calculated with respect to a catchment with a square shape that is equivalent to the real one and does not consider its actual shape, which can be elongated.	
Horton ratios	Horton–Strahler numbers	HS_num_u	–	Four sets of $u = 3$ vectors (each corresponding to a Horton order) containing, respectively, the number of streams of a given order [–], the average length of the streams of a given order [km], the average contributing area for each order [km ²] and the average slope of the streams of each order [%]. Slopes are calculated as the ratio of the difference in elevation between the ends of the segment to its length.
		HS_length_u	–	
		HS_area_u	–	
		HS_slope_u	–	
	Area ratio	R_a	–	Ratio of the average area drained by streams of a given order $u + 1$ and streams of order u .
	Bifurcation ratio	R_b	–	Ratio of the number of stream branches of a given order u to the number of stream branches of the next order, $u + 1$.
Length ratio	R_l	–	Ratio of the average stream lengths of two adjacent orders u and $u + 1$.	
Slope ratio	R_s	–	Ratio of the average stream slopes of two adjacent orders u and $u + 1$.	
Streamflow network	Total stream length	TSL	km	Total length of the river network, which is obtained by summing the lengths of all its segments.
	Drainage density	drain_dens	km ⁻¹	Ratio between the total stream length and the catchment area.
	Length of main channel	LMC	km	Length of the longest succession of segments that connect a source to the outlet of the catchment.
	Length of the longest drainage path	LLDP	km	Path included between the outlet and the furthest point from it that is placed on the catchment boundary and identified by following the drainage directions. For most of its length, the longest drainage path overlaps the main channel.
	Topological diameter	topo_d	–	Number of confluences found along the main channel.
	Mean hillslope length	MHL	km	Average of the distances (measured following the drainage directions) of all the pixels that do not belong to the hydrographic network, starting from the first pixel of the hydrographic network into which they drain.
	Mean slope of the longest drainage path	LLDP_slope	%	Calculated as $\text{LLDP_slope} = \frac{1}{\text{topo_d}} \sum \frac{\Delta z_i}{L_i} \cdot 100, \quad (1)$ where topo_d is the topological diameter, L_i is the length of the i th segment into which LLDP is divided and Δz_i is the corresponding elevation difference.
Shape factor and amplitude function	Shape factor	shape_f	–	Ratio of the catchment area to the square of the length of the main channel.
	Elongation ratio	elong_r	–	Ratio of the diameter of a circle of the same area as the catchment to the maximum drainage path length.
	Circularity ratio	circ_r	–	Ratio of the catchment area to the area of a circle with the same circumference as the perimeter of the catchment.
	Compactness coefficient	c_c	–	Ratio of the catchment perimeter to the diameter of a circle with the same area as the catchment.
	Width function characteristics	width_mean width_var width_skw width_kur width_x	–	Frequency distribution of the distances (along the drainage path) of the cells of the catchment to the outlet. The first four statistical moments (mean, variance, skewness and kurtosis) of this function were calculated, as was the percentiles vector containing the distances to the outlet that include pixel percentages of $x = 5 \%, 10 \%, 15 \%, 30 \%, 40 \%, 50 \%, 60 \%, 70 \%, 85 \%$ and 95% .

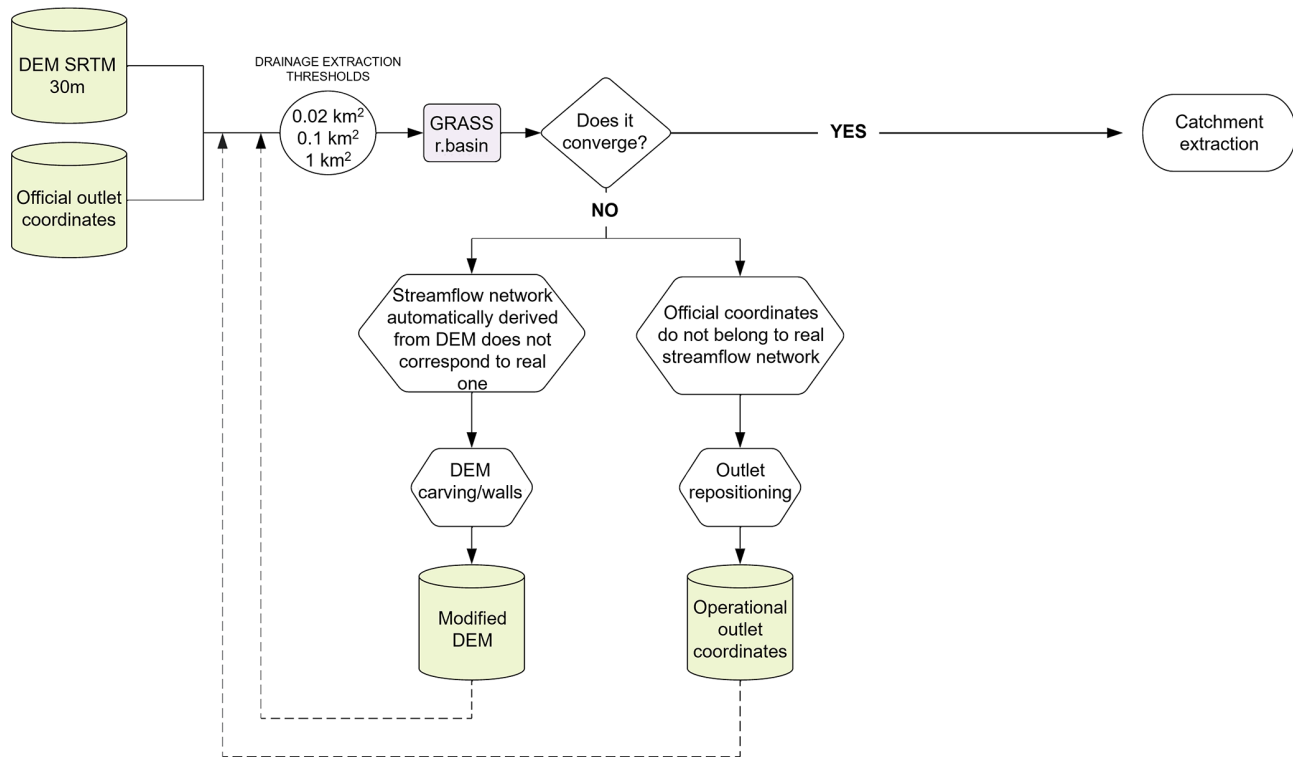


Figure 1. Visual representation of the workflow adopted to delineate catchment boundaries.

runoff formation. We believe that the small number of descriptors included in previous works could be insufficient to fully characterize areas with complex river networks. Here we stress the need to carefully characterize the catchments from a geomorphological point of view too, especially when a high percentage of the catchments considered are located in mountainous areas.

To further highlight the importance of this latter point, we underline that in the Italian context, which is the focus of this work, non-standard geomorphological parameters have already been demonstrated to be significant in several applications of statistical hydrology, such as the regionalization of the flood frequency curves for NW Italy developed by Laio et al. (2011). In that case, the proposed regression models consider geomorphological parameters like the area, the mean elevation, the length of the longest drainage path, the length of the orientation vector and the catchment outlet coordinates as covariates. In a revised version of the methodology (Ganora et al., 2014), the minimum elevation and the shape factor were also significant covariates. In Ganora et al. (2023), a non-dimensional flood reduction function to estimate design hydrographs for ungauged catchments was obtained with a multiple linear regression model that included the longest drainage path length and the slope of the catchment, the average catchment elevation, and the fourth statistical moment of the width function.

Moving to a broader context, the literature suggests that among the factors that govern the catchment's hydrological response during the process of rainfall–runoff transformation, the catchment's geomorphological features have been well represented (see, e.g., Nagy et al., 2021; Ravazzani et al., 2019). The synthesis of the basin response function from physical basin characteristics becomes crucial in ungauged catchments (Singh et al., 2013). Finally, most of the formulations available in the literature to estimate the response time when no runoff data are available typically contain a characteristic length and the slope of the catchment or the main channel (e.g., Chow, 1962; Kirpich, 1940; Sheridan, 1994), which are not readily available from the other above-mentioned databases.

The catchments included in the FOCA dataset cover a wide variety of morphological features and a considerable elevation range (Fig. 2a). Figure 2 also shows the empirical cumulative distribution functions (ECDFs) of some key catchment attributes: the area (Fig. 2b) and the mean catchment elevation (Fig. 2c). Figure 2 also reveals the limited number of catchments with outlets close to the sea. However, there is a reason behind this absence: the 631 chosen catchments that we considered are those for which peak or daily discharges are available and are therefore all included in the *Catalogo delle Piene dei Corsi d'acqua Italiani*. This implies that all the gray areas are ungauged; they could be included in the

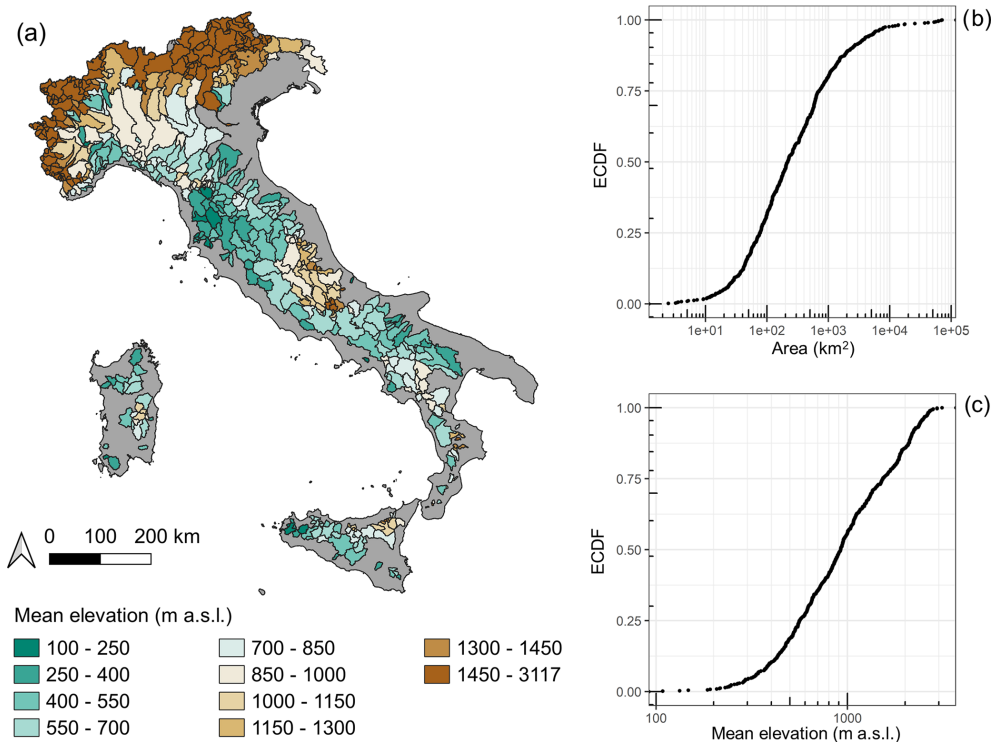


Figure 2. Mean elevation of the 631 catchments (a) and empirical cumulative distribution functions (ECDFs) of some key catchment attributes: area (b) and mean elevation (c). Administrative boundaries: GADM v3.6.

dataset but only if gauging stations were to be installed in the near future.

3.4 Quality check of the geomorphic data

To provide a robust set of catchment features, we have checked the consistency between this new dataset and the previously published one (Claps et al., 2020a, b, c) in terms of catchment areas, which were mostly based on the values published in the hydrologic yearbooks and in *Pubblicazione no. 17*. In order to set the level of accuracy of the procedures that we used, a 10 % maximum deviation of the difference (positive or negative) was considered acceptable. All the catchments for which the discrepancy is greater than 10 % have been individually re-examined, as described in Sect. 3.2. In some cases, the discrepancy was found to derive from a small shift of the catchment outlets upstream or downstream of a confluence, thus including or excluding relevant sub-catchments.

Another important check has been made: the difference between the length of the main channel (LMC) of each catchment and the length of the longest drainage path (LLDP) was analyzed. Even though these two attributes are well known to have a strong relationship, this connection was not always found in our data: for some catchments, the two lengths were significantly different, even after a thorough manual check. A threshold of a 3 km difference (corresponding to the value

that selects the upper 5 % of the catchments with marked differences) was used to identify those catchments for which the difference in length between LMC and LLDP should be further investigated. For those catchments, manual inspections highlighted a drawback in the GIS procedure that produced LMC and LLDP measurements with unrealistic discrepancies. Two different cases were observed. In one case, the resulting main channel shapefile consisted of a polygonal chain made up of multiple features that needed to be merged into one. In the other case, multiple LLDPs that differed from each other by no more than 100 m (2–3 pixels) were identified for the same catchment. In the latter case, one LLDP was manually chosen and the other ones were removed. We also observed that the two situations could occur simultaneously.

To get a general feel for the consistency of these two attributes, some scaling laws for the drainage network can be employed. One of the best-known is Hack's law (Hack, 1957; Eq. 3), which reads

$$\text{LLDP} = \alpha \text{ area}^\beta, \quad (2)$$

where the coefficients α and β vary depending on the study area. A log-log scatter plot between main channel length and basin area for all 631 catchments is displayed in Fig. 3, where panels (a) and (b) refer to the results before and after the quality control, respectively. By means of the abovementioned comparisons, we were able to double-check the con-

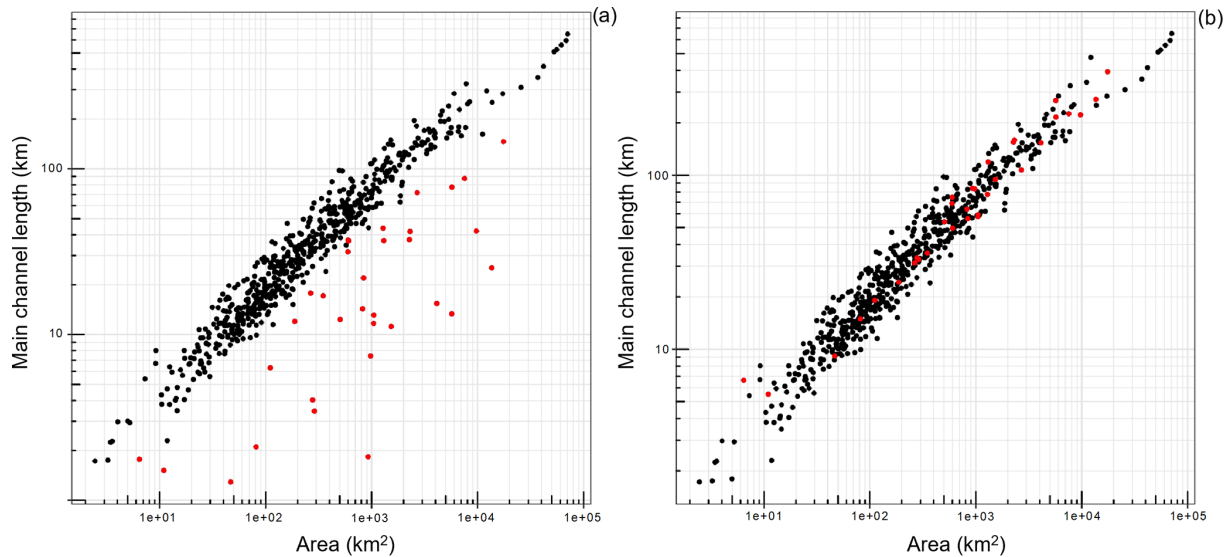


Figure 3. Scaling law between main channel length and basin area for the 631 catchments before (a) and after (b) the quality-control procedure for LLDP and LMC.

sistency between the two attributes; we found errors of up about 20 km, which we checked and corrected.

4 Data and catchment attributes concerning the soil, vegetation and climate

This section describes the work done and the datasets used to evaluate area-averaged catchment attributes concerning the soil, vegetation, climate and related extremes (for both rainfall and discharges).

For all the attributes, a spatial average value is provided, along with an additional assessment of the spatial variability in terms of the (spatial) coefficient of variation. On the other hand, the temporal variability is not considered here. The only exceptions are the coefficient of variation of the NDVI (normalized difference vegetation index) and the coefficient of variation of the rainfall regimes, which is introduced later and is described in Table 4. The spatial coefficient of variation can thus be considered a proxy for the uncertainty associated with the use of a mean value for the entire catchment. This statistic is additional information that is not usually reported in the CAMELS datasets, but it represents added value here, especially for a nation with such a complex geomorphology and climate as Italy.

All the attributes that will be described hereafter are affected by different sources of uncertainty that may or may not be quantified. To avoid the negative influence that resampling of the original data (available at different spatial resolutions) to a standard specific resolution could have on the results, in this work we used input data with an adequate resolution in relation to the type of information, so we avoided resampling. Moreover, when dealing with a raster obtained

through the spatial interpolation of at-site values, the spatialization method could introduce a certain level of uncertainty at the local scale, which could be quantified in pixels where a measurement was present. For the input data that we used, information about measurement or interpolation uncertainty is not always available, since we used several raster maps produced by other research groups. This is the case for land cover, NDVI and climatological data. Information about the uncertainty of the soil data that we used is available in Hengl et al. (2017). When referring to datasets produced by the authors involved in the creation of FOCA (discharge and rainfall extremes; see Sects. 4.2 and 4.3), some additional information is provided.

4.1 Soil, land cover and NDVI catchment attributes

In this subsection, different rasterized information is analyzed to provide area-averaged values. The spatial resolutions of the raw data range from 100 to 1000 m, and the data are not resampled at the same resolution before computing the average values.

Soil descriptors included in the FOCA dataset can provide information connected to an area-averaged estimate of the soil permeability conditions. These descriptors are the mean areal values of the curve number (Soil Conservation Service, 1972) and of the saturated hydraulic conductivity.

The curve number is an empirical parameter used to evaluate the portion of the total rainfall that becomes the net rainfall in a flood event. The values used in this study were taken from a national-scale cartography produced by Carrero (2004) at 250 m resolution, consistent with the work of Ganora et al. (2013). Since this dataset is currently not available online, it was also included in the dataset in the form of

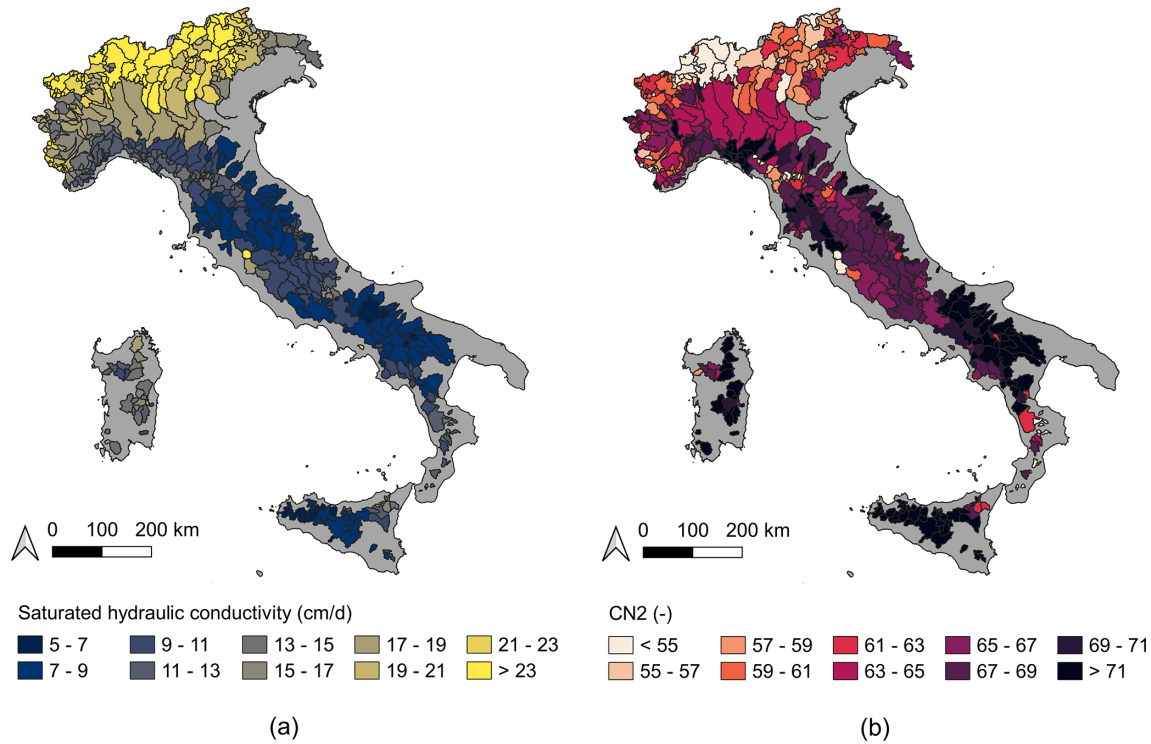


Figure 4. Mean areal values of the saturated hydraulic conductivity (a) and the curve number in average conditions, i.e., CN2 (b). Administrative boundaries: GADM v3.6.

Table 2. List of soil attributes.

Attribute sub-category	Attribute	Notation	Units	Description
Soil	Curve number	CN1, CN1_cv, CN2, CN2_cv, CN3, CN3_cv	–	Empirical parameter developed by the Soil Conservation Service (1972) and used to predict the direct runoff. Its value is between 0 and 100. According to the antecedent moisture conditions, i.e., the preceding wetness of the soil, CN is divided into three classes, namely CN1, CN2 and CN3. CN2 is the average condition, while CN1 and CN3 represent the lowest (dry soil) and highest (saturated soil) runoff potentials, respectively.
	Saturated hydraulic conductivity	k	cm d ⁻¹	Computed from the sand and clay contents of SoilGrids maps at 250 m resolution as follows: $k_s = 24e^{\left[\frac{12.012 - 7.55 \cdot 10^{-2} s + \frac{(-3.895 + 3.671 \cdot 10^{-2} s - 0.1103c + 8.7546 \cdot 10^{-4} c^2)}{0.332 - 7.251 \cdot 10^{-4} s + 0.1276 \log(c)}} \right]}$ (3) where <i>s</i> is the sand content (%) and <i>c</i> is the clay content (%).

a raster map to grant replicability and thus allow researchers to perform the same study or extend this one to other catchments. Three types of curve numbers are defined, which relate to three different antecedent wetness conditions – dry, average and wet, respectively – of the soils. For each curve number type, we provided the mean value and the (spatial) coefficient of variation.

To derive a feature that approximates soil permeability features, we started from soil texture fraction maps derived from SoilGrids (Hengl et al., 2017; available at <https://soilgrids.org/>, last access: 23 October 2023). These cartographies map the spatial distribution of soil properties across the globe at 250 m spatial resolution and at seven standard depths ranging from 0 to 200 cm. SoilGrids maps are based on over 230 000

soil profile observations from the WoSIS (World Soil Information Service) database (Batjes, 2009). Soil texture information was derived from these maps after averaging across the first 30 cm of depth, a value consistent with the hydrological purposes of this work. Based on the sand and clay contents, we derived the saturated hydraulic conductivity using the pedo-transfer function proposed by Saxton et al. (1986; Eq. 3).

The seven soil descriptors (six related to the curve number and one related to the saturated hydraulic conductivity) that we considered are listed in Table 2.

Figure 4 provides a snapshot map of the catchment-averaged soil parameter values. The resulting spatial distribution of the saturated hydraulic conductivity (Fig. 4a) illus-

trates that the soil properties vary significantly upon moving from the north to the south of Italy and reflects the high clay content characterizing the soils of the Apennine basins. The same level of difference is not visible in the CN2 values (Fig. 4b), essentially because they reflect the geological and land-use (not soil) information.

Land cover characteristics at 100 m resolution are extracted from the 44 classes of the third level of CORINE Land Cover 2018 (available at <https://land.copernicus.eu/>, last access: 23 October 2023). In particular, five land indices obtained by merging similar classes are considered here. The five land indices are outlined in Table 3. A map of one of the classes is provided in Fig. 5a: it displays an overview of the percentage of urbanized areas, thus providing some information related to the relevance of anthropogenic settlements within the catchments. As expected, the vegetation coverage is lower in high-elevation and cold regions (i.e., the Alps), which are also regions with low percentages of urbanized areas.

We also computed multi-temporal indicators of the NDVI (normalized difference vegetation index) maps, whose data are provided by the Copernicus Land Monitoring Service (available at <https://land.copernicus.eu/>, last access: 23 October 2023). The NDVI is an index that shows whether the area under investigation contains live green vegetation and shows its overall health. We used the values of the Long Term Statistics (LTS) NDVI V3.0.1 of the Copernicus service (with a 1 km spatial resolution), which are mean NDVI observations over the period 1999–2019 for each of the 36 periods of 10 days per year, resulting in 36 raster maps. These maps were used to compute the mean NDVI value, the (temporal) coefficient of variation of the NDVI and the spatio-temporal mean NDVI regime. The NDVI regime is a graphical representation of the multi-temporal mean calculated over 36 time intervals, each spanning 10 d. To synthetically characterize the latter, a Fourier series representation was used, which allows the shape of the regime to be described with fewer parameters (four in total) than the 36 average values (each for a different 10 d period) from which it is composed. A more detailed description of the four parameters that describe the coefficients of the Fourier series that represent the NDVI regime is reported in Appendix A.

After this data preparation, catchment boundaries were used to extract a total of 11 land cover and NDVI attributes, as listed in Table 3. The mean NDVI is mapped in Fig. 5b to provide some insight into the variation of the mean greenness of the biomes.

4.2 Climatological catchment attributes

State-of-the-art national-scale datasets at 1 km resolution have been used for the evaluation of several climatological attributes, as described below.

Mean monthly precipitation information is taken from the BIGBANG (Bilancio Idrologico GIS BAsed a scala

Nazionale su Griglia regolare; Braca et al., 2021) 4.0 dataset, which covers the 1951–2019 period. BIGBANG is obtained by means of the spatial interpolation of rain gauge measurements at 1 km resolution, and integrates, only over limited areas and for certain years, the spatial interpolation produced by ARCIS (Archivio Climatologico per l'Italia Centro Settentrionale; Pavan et al., 2019). Mean monthly temperature data are also derived from this dataset. Both mean monthly precipitation depths and mean monthly temperature data are processed to compute the mean coefficients of the Fourier series that approximate the precipitation and temperature regimes (four coefficients for rainfall and four for temperature; see Appendix A).

This dataset is also used to compute the mean annual precipitation (MAP) and the mean annual temperature (MAT).

Catchment boundaries were used to clip the abovementioned precipitation and temperature maps and obtain spatial averages for the 14 climatological attributes listed in Table 4.

Upon presenting the results, one can notice in Fig. 6a that the mean annual precipitation (MAP) is higher in mountainous regions, with the largest values seen in the Alps. The same areas also show a low mean annual temperature (MAT; Fig. 6b).

4.3 Extreme rainfall catchment attributes

With respect to other national databases, like the LamaHCE dataset (Klingler et al., 2021), one of the main novelties of this work is the introduction of information related to parameters of the sub-daily precipitation extremes. This inclusion was made possible thanks to the availability of a rich collection of in situ data that are characterized by a much greater accuracy when capturing extremes compared to reanalysis data like the ERA5 and ERA5-Land datasets (Muñoz Sabater, 2019) widely used in the creation of other datasets. The annual maximum rainfall depths used to derive spatial extreme rainfall statistics are obtained from the Improved Italian – Rainfall Extreme Dataset (I²-RED; Mazzoglio et al., 2020). This dataset consists of official and quality-controlled short-duration (1, 3, 6, 12 and 24 h) annual maximum rainfall depths recorded by more than 5200 rain gauges across Italy between 1916 and 2019. The stations were subjected to a quality-control procedure to correct errors in the plano-altimetric positions, duplicates and incorrect rainfall measurements (Mazzoglio et al., 2020). This FOCA dataset represents the first national-scale collection of mean extreme rainfall catchment attributes. Thanks to FOCA, it is now possible to perform simple regional and national hydrological studies without the need to retrieve information from 21 different agencies, as mentioned in the “Introduction”. Due to the complex data policy that regulates the data collected by the hydrological agencies (most of them only provide the data free of charge for research purposes, while people interested in using them for commercial purposes are requested to pay), we were not allowed to include

Table 3. List of land-use and NDVI attributes.

Attribute sub-category	Attribute	Notation	Units	Description
Land cover	CORINE Land Cover 1	clc1	%	Percentage of the catchment area that corresponds to continuous and discontinuous urbanized areas (CORINE classes 111, 112).
	CORINE Land Cover 2	clc2	%	Percentage of the catchment area that corresponds to woods (311, 312, 313), arboreal vegetation, shrub vegetation and bushes (CORINE classes 324, 323, 321, 322).
	CORINE Land Cover 3	clc3	%	Percentage of the catchment area that corresponds to herbaceous vegetation, meadow pasture, special crops, olive groves, vineyards and arable land (CORINE classes 231, 222, 223, 221, 211, 241, 243, 242, 142).
	CORINE Land Cover 4	clc4	%	Percentage of the catchment area that corresponds to non-vegetated areas (331, 333, 332, 334), mining areas, landfills, construction sites (CORINE classes 131, 133), industrial and commercial areas, and communication networks (CORINE classes 121, 122, 123, 124).
	CORINE Land Cover 5	clc5	%	Percentage of the catchment area that corresponds to humid areas (CORINE classes 411, 512, 521).
NDVI	NDVI	NDVI NDVI_cv	–	Indicator of the greenness of the biomes measured by satellite, which has a value of between 0 and 1. It is defined as $NDVI = \frac{REF_{nir} - REF_{red}}{REF_{nir} + REF_{red}}, \quad (4)$ where REF_{nir} and REF_{red} are the spectral reflectances measured in the near-infrared and red wavebands, respectively. Its mean value and (temporal) coefficient of variation were computed.
	B1, B2, C1, C2	B1_NDVI, B2_NDVI, C1_NDVI, C2_NDVI	–	Mean values of the coefficients of the Fourier series representation of NDVI (see Appendix A).

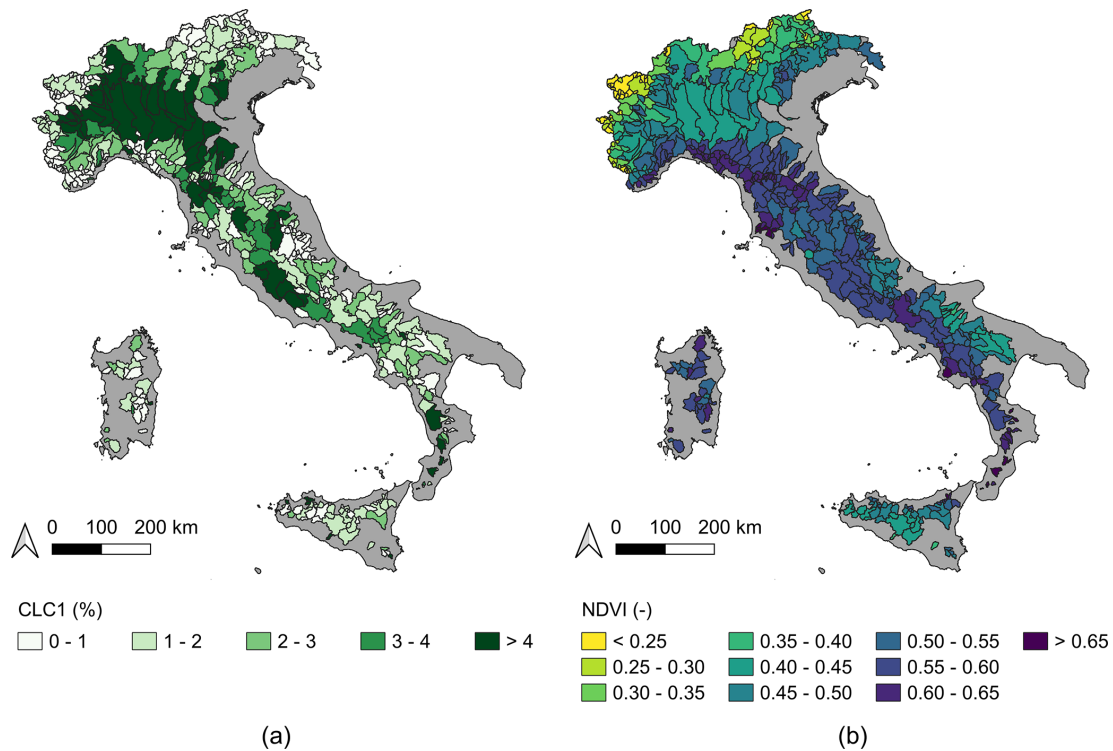
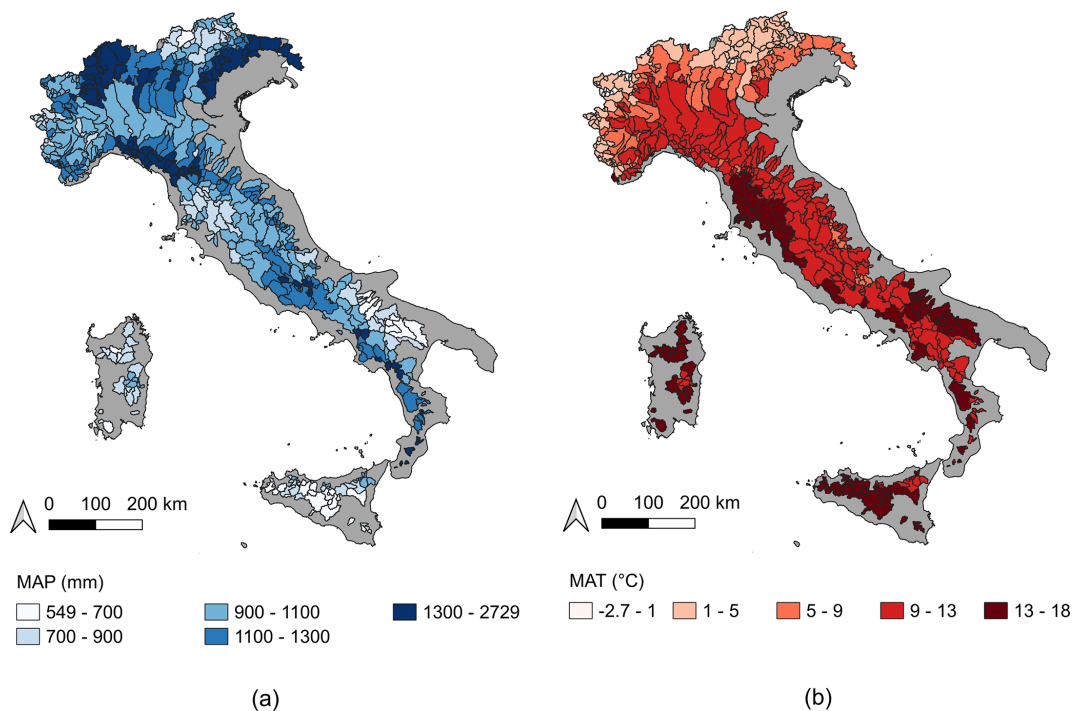


Figure 5. Spatial variability of the percentage of clc1 (a) and mean areal NDVI (b). Administrative boundaries: GADM v3.6.

Table 4. List of climatological attributes.

Attribute	Notation	Units	Description
Mean annual precipitation	MAP MAP_cv	mm	Spatial mean and coefficient of variation of the total mean annual precipitation (Braca et al., 2021).
B1, B2, C1, C2	B1_rain, B2_rain, C1_rain, C2_rain	–	Mean values of the coefficients of the Fourier series representation of the precipitation regime computed from the mean monthly precipitation (see Appendix A).
Coefficient of variation of the rainfall regimes	cv_rain	–	Temporal coefficient of variation calculated from the monthly mean precipitation depths derived from Braca et al. (2021).
Time step between the maximum and minimum mean monthly rainfall	seas_prec	–	Number of months between the occurrence of the absolute annual maximum precipitation and the subsequent absolute minimum precipitation.
Mean annual temperature	MAT MAT_cv	°C	Spatial mean and coefficient of variation of the mean annual temperature computed using Braca et al. (2021).
B1, B2, C1, C2	B1_temp, B2_temp, C1_temp, C2_temp	–	Mean values of the coefficients of the Fourier series representation of the temperature regimes (see Appendix A).

**Figure 6.** Spatial representations of the mean annual precipitation (a) and mean annual temperature (b). Administrative boundaries: GADM v3.6.

the annual maxima time series in the FOCA dataset. However, nationwide maps of the parameters allowing the computation of the IDF (intensity–duration–frequency) curves have been prepared and made freely available within the FOCA dataset.

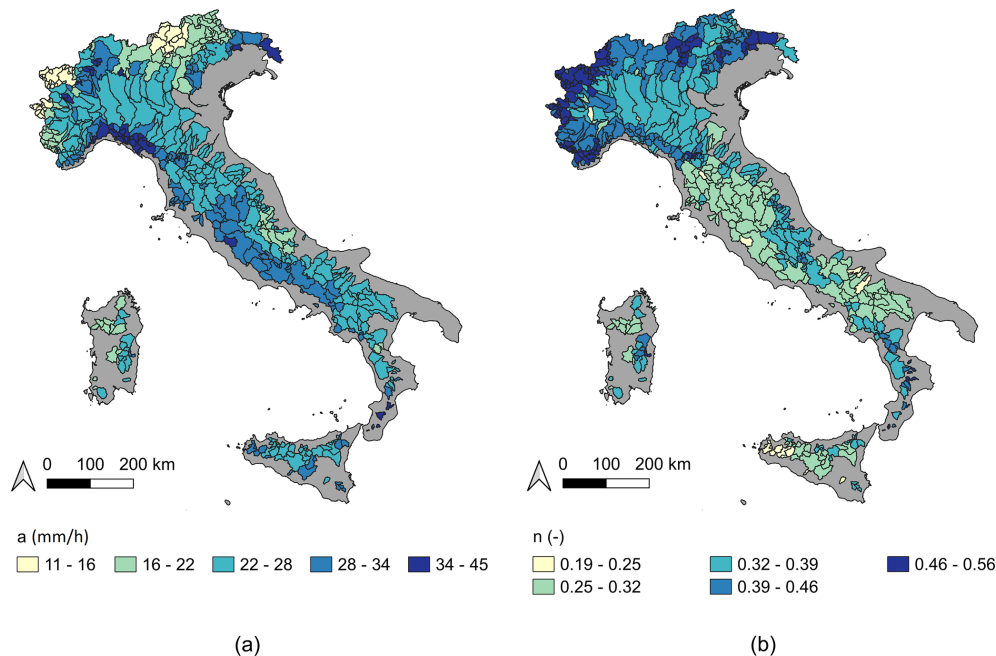
Rainfall data and related statistics are processed at 250 m resolution with the R function *autokrige* (Hiemstra and

Skoien, 2023). This performs automatic ordinary kriging using the variogram that best fits the data, which is automatically generated by the R function *autofitVariogram*. The rainfall statistics obtained with this procedure are

- the scale factor a and the scaling exponent n of the average intensity–duration (ID) curves, which are obtained

Table 5. List of rainfall attributes.

Attribute	Notation	Units	Description
a	a a_{cv}	mm h^{-1}	Scale factor of the IDF curve. The mean value and (spatial) coefficient of variation were computed.
n	n n_{cv}	–	Scaling exponent of the IDF curve. The mean value and (spatial) coefficient of variation were computed.
L-CV d h	LCV_{dh} $LCV_{dh_{cv}}$	–	Coefficient of L-variation for $d = 1, 3, 6, 12$ and 24 h duration. The mean value and (spatial) coefficient of variation were computed.
L-CA d h	LCA_{dh} $LCA_{dh_{cv}}$	–	Coefficient of L-skewness for $d = 1, 3, 6, 12$ and 24 h duration. The mean value and (spatial) coefficient of variation were computed.

**Figure 7.** Spatial representations of the spatial averages of parameter a (a) and parameter n (b) of the ID curves. Administrative boundaries: GADM v3.6.

by linear regression of the logarithm of the average rainfall depth h_d for the durations of 1 to 24 h against the logarithm of the duration, where

$$h_d = a \cdot d^n \quad (5)$$

- the coefficient of L-variation (L-CV) for the durations of 1 to 24 h, which is evaluated with Eq. (6) of Laio et al. (2011)
- coefficient of L-skewness (L-CA) for the durations of 1 to 24 h, which is evaluated with Eq. (7) of Laio et al. (2011).

Time series with at least 10 years of data were used to evaluate the parameters a and n , while series with 20 and 30 years of data were used for L-CV and L-CA, respectively.

The different record lengths were selected because higher-order statistics cannot be evaluated from short time series (Koutsoyiannis, 2019). These maps represent the first attempt to reconstruct and represent updated extreme rainfall statistics across Italy following similar releases for Switzerland (i.e., the *Hydrological atlas of Switzerland*, available at <https://hydrologicalatlas.ch/>, last access: 23 October 2023), Austria (i.e., the *Hydrological atlas of Austria*; Fürst et al., 2009), Germany (i.e., the KOSTRA-DWD or *Coordinated heavy precipitation regionalization and evaluation of the DWD*, available at https://www.dwd.de/DE/leistungen/kostra_dwd_rasterwerte/kostra_dwd_rasterwerte.html, last access: 23 October 2023), the United States (i.e., NOAA Atlas 14, available at https://www.weather.gov/owl/hdsc_currentpf, last access: 23 October 2023 and at <https://hdsc.nws.noaa.gov/hdsc/pfds/>, last access: 23 October 2023), Hawaii (i.e.,

the *Rainfall atlas of Hawai'i*, available at <http://rainfall.geography.hawaii.edu/>, last access: 23 October 2023) and Canada (CSAGroup, 2019).

Table 5 summarizes the characteristics of the attributes computed and averaged over the 631 catchments.

A discussion of the overall spatial and temporal variability of the rainfall extremes is available in Mazzoglio et al. (2020, 2022a, b, 2023). What is useful to comment on here is the result of the spatial averaging of the parameters a and n at the catchment scale (Fig. 7a and b). The parameter a can be combined with n to obtain the mean rainfall for different durations (1 to 24 h). The first one represents the mean rainfall extremes of 1 h duration: the higher the value of a , the higher the intensity of short-duration rainfall extremes over the catchments. This parameter is particularly relevant for small catchments with a characteristic time lag of the order of 1 h. If Fig. 7 is compared to Fig. 2, we can see a modest correlation between a and elevation (the reverse orographic effect; Avanzi et al., 2015; Mazzoglio et al., 2022a, 2023). One can also notice that the regions where the most severe short-duration rainfall events occur (Fig. 7a) are different from those that present a higher MAP (Fig. 6a), which is a characteristic of the Mediterranean climate (Mazzoglio et al., 2022a).

In terms of the assessment of the uncertainty associated with the spatial average values, Fig. 8 provides interesting insights, as it shows the standard deviation of the kriging estimates of the parameter a in each individual pixel. The figure shows an almost uniform spatial distribution of the interpolation uncertainty, with higher values found near the coastline (due to the lower station network density) and lower values found where the station density is high. Thus, this map not only serves to show the interpolation uncertainty but also provides information on the rain gauge density throughout the country.

4.4 Peak and daily discharge collections

As outlined in Sect. 2, the peak and daily discharges from each gauging station come from Claps et al. (2020a, b, c), which is the most up to date systematic flood collection in Italy. Until the 1970s, the data reflect most of the content of *Pubblicazione no. 17* of the SIMN (Servizio Idrografico, 1980). After that, a consistent integration was carried out by merging data from different sources (such as Bencivenga et al., 2011; Barbero et al., 2012; Hall et al., 2015; Brath et al., 2017; Settore Idrologico e Geologico Regionale, 2022; ARPA Lombardia Sistema Informativo Idrologico, 2023) using the same database set up for the CUBIST project (Claps et al., 2008).

The provision of that information in this paper (in terms of both the time series and of their statistics, like the mean values of the flood peaks, Q_p) offers the opportunity to fully characterize the climatology of the extremes for the catchments upstream of the gauging stations. The uncertainty of

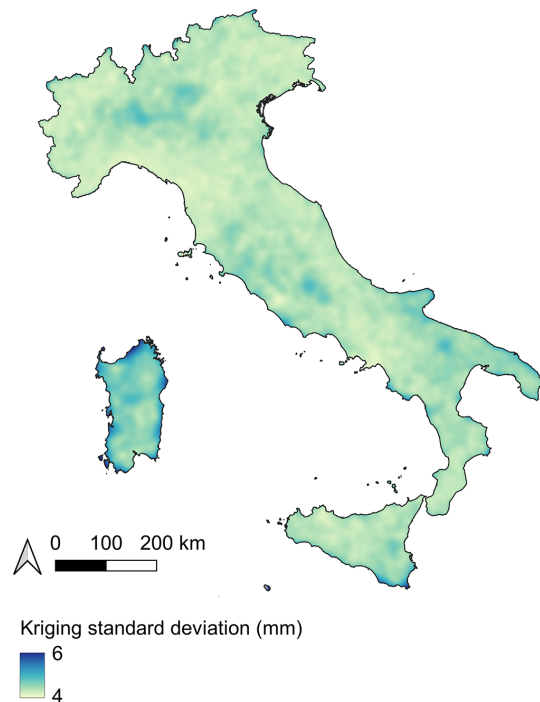


Figure 8. Standard deviation of the kriging estimate of the parameter a . Administrative boundaries: GADM v3.6.

observed discharge values cannot be reported since it is a piece of information that is not available to the public. In essence, before their publication in the hydrological yearbooks, these data were validated through the application of the corresponding rating curves for every year of measurement. However, it is not possible to rate the uncertainty of the level-to-discharge conversions, which were made decades ago. This would involve accounting for very site-specific technical issues that can be dealt with only in strict cooperation with the offices where the peak discharge values were validated, and it would need to be done for each and every measurement section.

The spatial variability of the mean annual flood is shown in Fig. 9a, in which one can recognize the presence of nested sub-catchments resulting from the presence of different measurement stations located along the river path. This situation is particularly evident in the northwest of Italy, where different small mountainous sub-catchments of the Po basin are highlighted, while the Po basin itself is the partially visible catchment depicted in black. To get an idea of the lengths of the discharge time series available, one can refer to Fig. 9b: the different dot colors show that most of the longer time series (> 50 years) are generally available in North and Central Italy. Moreover, while the reference period covered by the dataset dates from 1911 to 2016, the time series are fragmented and characterized by different time coverages and thus do not characterize the entire interval (Fig. 10). The decreasing number of time series in recent decades is due to

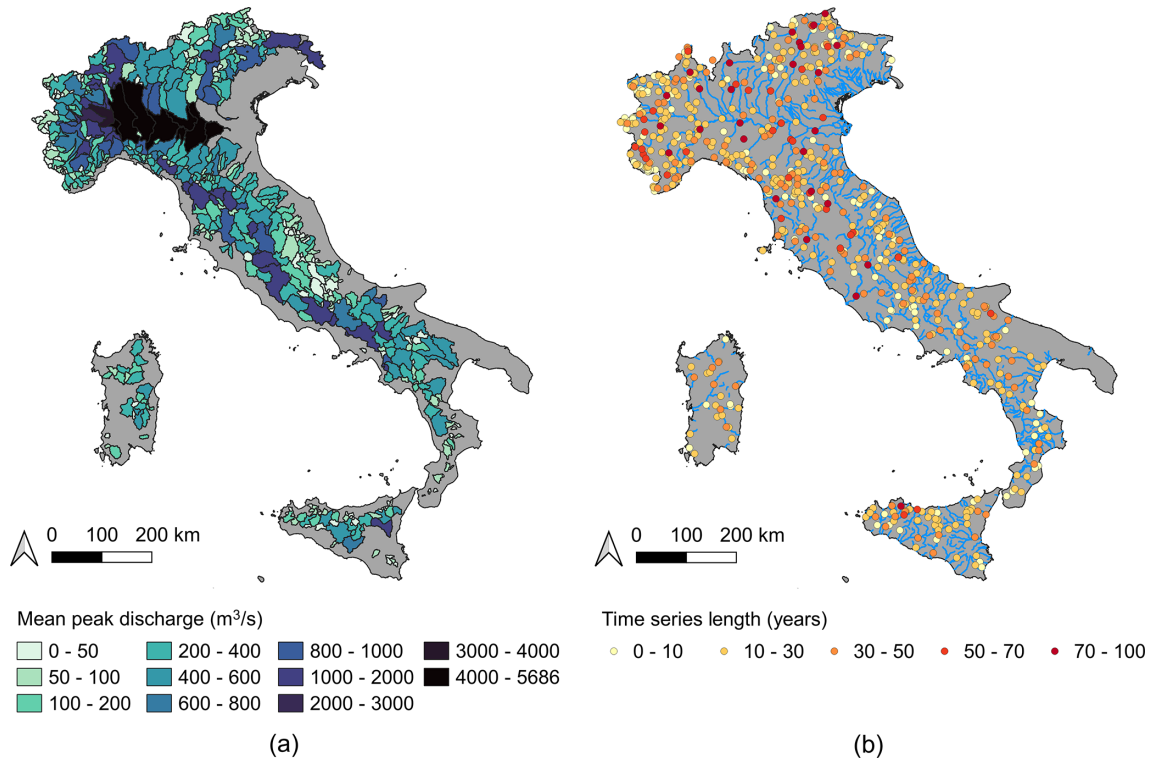


Figure 9. Spatial distributions of the mean peak discharge (a) and discharge time series length (b). Administrative boundaries: GADM v3.6.

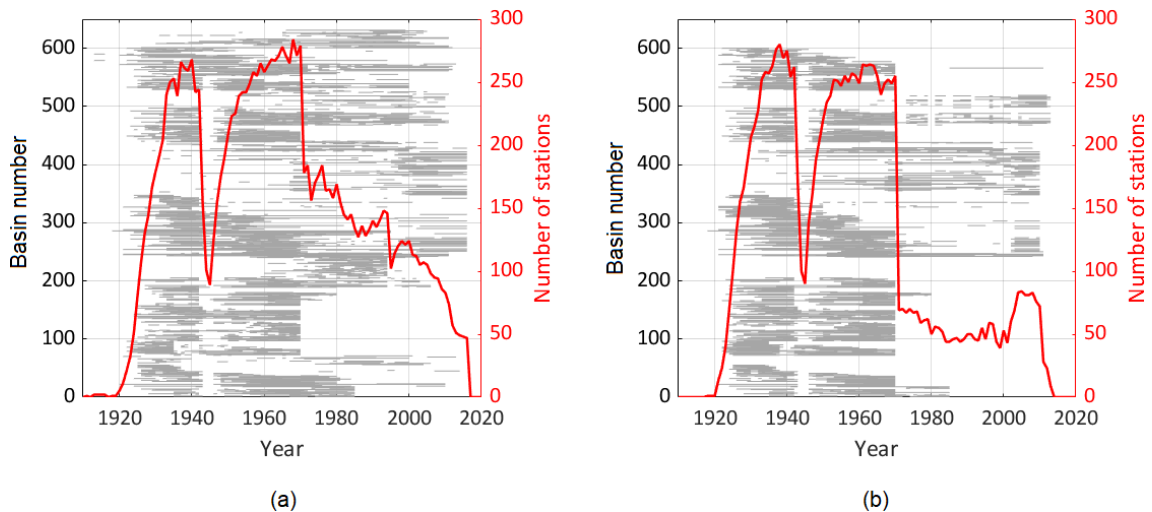


Figure 10. Temporal availability of the peak (a) and daily (b) discharge. Each gray line represents a time series; in these time series, gray identifies available records while white indicates missing data. The bold red line identifies the total number of measurements available for each year.

the lack of quality-controlled discharge data published by the regional hydrological agencies, which is mostly due to the unavailability of updated rating curves. In most cases, the regional agencies are now only publishing water levels, while information about discharges is often missing.

The catchments cover a wide variety of morphological features. Figure 11 shows the empirical cumulative distribution function (ECDF) of the specific mean annual flood (i.e., the ratio between the mean peak discharge Q_p and the catchment area; Fig. 11b). The relationship between the area, the specific mean annual flood and the mean catchment elevation

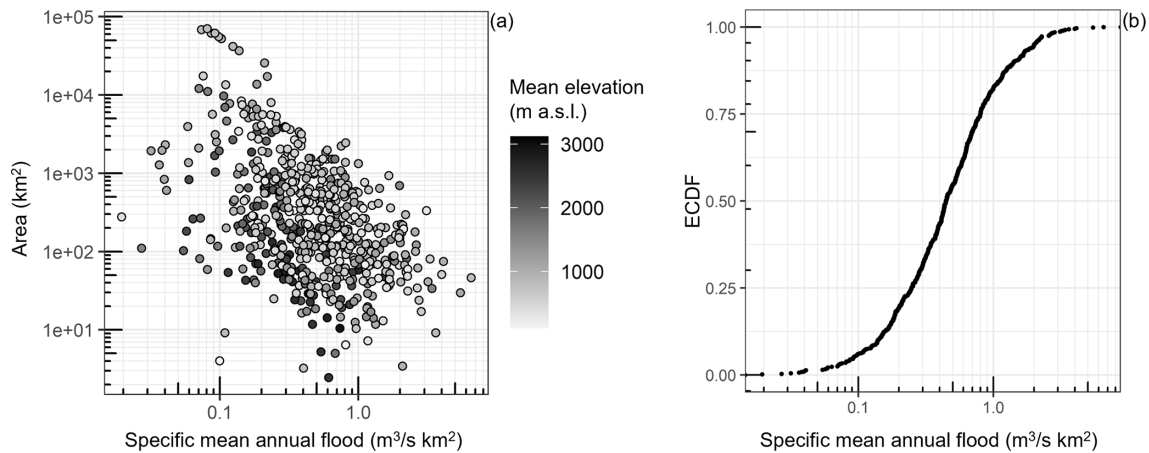


Figure 11. Relationship between catchment area, specific peak discharge and mean elevation (a) and the empirical cumulative distribution function (ECDF) of the specific mean annual flood (b).

is represented in Fig. 11, which shows how the Italian catchments are characterized by different hydrological regimes and a modest dependence of the mean elevation on the specific peak discharge. Catchments with a high mean elevation (from 2000 to more than 3000 m a.s.l.) cover a wide range of possible specific mean annual floods, but the interval is smaller than that characterizing catchments located at lower elevations.

5 Data availability

The dataset detailed in this paper is available at <https://doi.org/10.5281/zenodo.10446258> (Claps et al., 2023). It contains all the catchment boundaries and related catchment attributes described before. In case of future updates, to access the latest version of the database, readers can refer to <https://doi.org/10.5281/zenodo.8060736> (Claps et al., 2023) to download the most recent version.

6 Conclusions

In this work, we have presented FOCA (Italian FIOod and Catchment Atlas), a collection of attributes of 631 Italian river catchments derived from data sources that comply with the following criteria: (a) nationwide coverage; (b) consistency in data quality (i.e., no regional or local biases); and (c) adequate original resolution in relation to the type of information.

FOCA represents the most up to date collection of catchment boundaries and related attributes at the national scale for Italy. It summarizes features related to climate (discharge, rainfall and temperature), geomorphology, soil and land use. The dataset covers an overall area representative of most of the landscapes of Italy, including the mountainous ones (elevations ranging from 0 to 4800 m). One of the main novelties

with respect to other national-scale datasets is the inclusion of a rich set of geomorphological and extreme rainfall attributes. The second notable novelty is the inclusion, in the same dataset, of the most up to date information regarding extreme rainfall and discharge.

As mentioned before, Italy shows very wide variability in climate, land and morphological features, and this peculiarity emerges from this new dataset. Figure 2b shows that the FOCA dataset includes catchments with significantly different areas, with small-to-medium ones prevailing: 105 out of 631 catchments have an area of less than 50 km², and 300 out of 631 catchments have an area of less than 200 km². In addition, 279 out of 631 catchments have a mean elevation higher than 1000 m a.s.l., highlighting that, despite the existence of some large plains, Italy is a mountainous country. On the other hand, in the large catchments in the north of Italy, the spatial averaging operation over the catchment area becomes less significant, as it is evaluated over an area within which significantly different features coexist.

A key decision that we had to take while developing FOCA was whether to use local or global/quasi-global datasets or both. The use of global datasets facilitates the comparison of the results obtained in large-scale hydrology works, even when continental analyses are performed. However, local datasets are, without a doubt, characterized by higher-quality information. In this work, we decided to opt for the second approach, selecting the best possible data for each variable, prioritizing local information, and resorting to global data in only a few cases. The availability of digital information on the catchment boundaries allows users to evaluate the mean areal values of other variables, such as evapotranspiration and snow cover, using their own models or other large-scale reanalysis datasets, like ERA5 and ERA5-Land.

The FOCA dataset provides new opportunities to perform both regional and national-scale studies using catchment shapes and attributes extracted using a common frame-

work and subjected to the same quality-control procedure. Information about the algorithms used in this work are also reported in Sect. 3 to ensure replicability and allow the calculation of the attribute characteristics for any ungauged basin. Moreover, the database provides the opportunity to investigate how catchment attributes control river flows and allows the improvement of data-intensive investigations, such as performing predictions for ungauged catchments.

Appendix A

The rainfall, temperature and NDVI regimes are described using Fourier series, which allows the shape of the regime to be reconstructed using a reduced number of parameters compared to the 12 different monthly values (or the 36 values for 10 d periods in the case of the NDVI) used to describe it. According to Fourier's theorem, a wave of period τ and pulse $\omega = 2\pi/\tau$ can be described as

$$f(t) = A_0 + \sum_{i=1}^N (A_i \cos(i\omega t + \phi_i)), \quad (\text{A1})$$

where t is the time, A_0 is the mean of $f(t)$ over t , N is the number of harmonics, A_i is the harmonic amplitude and ϕ_i is the phase. In the case of two harmonics, Eq. (6) can be written as

$$\begin{aligned} f(t) &= A_0 + A_1 \cos\left(\frac{2\pi}{\tau}t + \phi_1\right) + A_2 \cos\left(\frac{4\pi}{\tau}t + \phi_2\right) \\ &= A_0 + A_1 \cos\left(\frac{2\pi}{\tau}t\right) \cdot \cos(\phi_1) - A_1 \sin\left(\frac{2\pi}{\tau}t\right) \\ &\quad \times \sin(\phi_1) + A_2 \cos\left(\frac{4\pi}{\tau}t\right) \cdot \cos(\phi_2) \\ &\quad - A_2 \sin\left(\frac{4\pi}{\tau}t\right) \cdot \sin(\phi_2). \end{aligned} \quad (\text{A2})$$

By separating the variables that depend on the time from those that are not time dependent, we obtain

$$f(t) = A_0 + B_1 \cdot X_1(t) + C_1 \cdot Y_1(t) + B_2 \cdot X_2(t) + C_2 \cdot Y_2(t). \quad (\text{A3})$$

In this analysis, we evaluated B_1 , B_2 , C_1 and C_2 with the ordinary least-squares method (see Eqs. 9 to 12).

$$B_1 = A_1 \cos(\phi_1) \quad (\text{A4})$$

$$B_2 = A_2 \cos(\phi_2) \quad (\text{A5})$$

$$C_1 = -A_1 \sin(\phi_1) \quad (\text{A6})$$

$$C_2 = -A_2 \sin(\phi_2) \quad (\text{A7})$$

The equation can be further simplified by assuming the mean A_0 to be null and solving it for $\tau = 2\pi$.

Author contributions. PC: conceptualization, funding acquisition, methodology, project administration, resources, supervision, writing – review and editing. GE: conceptualization, data curation,

formal analysis, investigation, methodology, software, validation, visualization, writing – original draft preparation, writing – review and editing. DG: conceptualization, methodology, writing – review and editing. PM: conceptualization, data curation, formal analysis, investigation, methodology, software, validation, visualization, writing – original draft preparation, writing – review and editing. IM: conceptualization, data curation, formal analysis, investigation, methodology, software, validation, visualization, writing – original draft preparation, writing – review and editing.

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