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Efficiency of closed loop geothermal heat pumps: a sensitivity analysis

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10 Abstract

11 Geothermal heat pumps are becoming more and more popular as the price of fossil fuels is 12 increasing and a strong reduction of anthropogenic CO₂ emissions is needed. The energy 13 performances of these plants are closely related to the thermal and hydrogeological properties of the 14 soil, but a proper design and installation also plays a crucial role. A set of flow and heat transport 15 simulations has been run to evaluate the impact of different parameters on the operation of a GHSP. 16 It is demonstrated that the BHE length is the most influential factor, that the heat carrier fluid also 17 plays a fundamental role, and that further improvements can be obtained by using pipe spacers and 18 highly conductive grouts. On the other hand, if the physical properties of the soil are not surveyed 19 properly, they represent a strong factor of uncertainty when modelling the operation of these plants. 20 The thermal conductivity of the soil has a prevailing importance and should be determined with in-21 situ tests (TRT), rather than assigning values from literature. When groundwater flow is present, the 22 advection should also be considered, due to its positive effect on the performances of BHEs; by 23 contrast, as little is currently known about thermal dispersion, relying on this transport mechanism 24 can lead to an excessively optimistic design.

25

26 Keywords:

27 Low-enthalpy geothermal energy, Borehole Heat Exchanger, Ground Source Heat Pump, Heat28 transport, Groundwater

29 **1. Introduction**

Ground Source Heat Pumps (GSHP) are space heating and cooling plants which exploit the soil as a 30 31 thermal source or sink, through the circulation of a heat carrier fluid in a closed pipe loop. Different 32 pipe arrangements are available, among which the most common is the Borehole Heat Exchanger, a 33 vertical pipe loop reaching depths of 50 to 200 m (Fig.1). Below a depth of a few meters from the 34 ground surface, the seasonal variation of the air temperature disappears due to the large thermal 35 inertia of the soil. Therefore, if compared to the air, the soil is a warmer source for heating during winter and a cooler sink for cooling during summer, and higher system efficiencies can therefore be 36 37 achieved compared to Air Source Heat Pumps.

38 GSHPs are rapidly spreading in Europe, China and USA, and have a great potential for energy, cost 39 and CO₂ emission saving [1]. About 100,000 low-enthalpy geothermal plants are installed every 40 year in Europe, mainly for new dwellings in Sweden, Germany and France [2, 3]. According to 41 Saner et al. [4], the use of GSHP in place of methane furnaces allows the CO₂ emissions to be 42 reduced by up to 84%, depending on the sources used for the production of electricity. From the 43 economic point of view, the geothermal heat pumps lead to a considerable reduction of the 44 maintenance costs and, although their installation is more expensive than the other heating and 45 cooling plants, the payback periods proved to be reasonable, i.e. less than 10 years [5-7].

Since the thermal exploitation of the soil induces a gradual temperature drift, an accurate heat transport modelling of soil and aquifer systems is essential for a correct design of GSHPs. Indeed, the efficiency of the heat pump is strongly influenced by the temperature of the heat carrier fluid, which in turns depends on the temperature of the surrounding soil. To estimate the thermal impact of BHEs and the working temperatures of the heat carrier fluid, different methods have been developed, which can be divided into analytical, semi-analytical and numerical.

52 The Kelvin infinite line source [8] and the infinite cylindrical source [9] are the simplest analytical 53 methods for estimating the thermal disturbance induced by a BHE, since they rely on the 54 assumption of a purely conductive and radial heat transport. Their main limitation is that of not 55 accounting for the vertical thermal gradient and fluxes [10] and for the heterogeneity of the heat 56 exchange over the length. Moreover, the advective and dispersive heat transport occurring in 57 aquifer systems is also neglected. Nevertheless, these analytical solutions are still widely used for 58 the interpretation of Thermal Response Tests [11], since they last for a short time (48÷72 h) and 59 therefore the vertical heat transport can be neglected. The subsurface flow and the seasonal changes 60 of groundwater levels can significantly alter the results of a TRT, as pointed out by Bozdağ et al. 61 [12]. To overcome this problem, Wagner et al. [13] recently developed a method for the 62 interpretation of TRTs in the presence of strong groundwater flow.

63 The semi-analytical method proposed by Eskilson [14] takes into account the finite length of the 64 exchanger and different BHE field layouts, but the advection and the dispersion are neglected. This 65 method is applied by two of the most popular BHE design software programmes, Earth Energy 66 Design [15] and GLHEPRO [16].

Analytical models which take into account the beneficial effects of groundwater flow [17], of the finite length of the BHE [18], and both them together [19] have been developed in the last few years, and they could be used in the future for the dimensioning of BHE fields.

Recently, numerical modelling has often been applied to the design of BHE fields. The finitedifference modelling software MODFLOW can be used coupled with the solute transport package MT3D (or MT3DMS) and by applying the analogy between heat and solute transport [20, 21], or with the specific heat transport package SEAWAT [22]. On the other hand, the finite element software FEFLOW includes a special package for the simulation of BHEs [23, 24] which is particularly suitable for non conventional BHE field layouts and for taking into account the thermal advection and dispersion in aquifer systems.

77

78 The heat transport simulation of GSHPs permits the assessment of their performances, which are 79 influenced by the properties of the exchanger and the thermo-hydrogeological parameters of the

80 soil. According to Chiasson et al. [25], groundwater flow significantly enhances the performances 81 of BHEs, and the Peclet number is a good indicator for whether advective transport needs to be 82 taken into account or neglected. Wang et al. [26] have developed a method to estimate the velocity 83 of groundwater movement measuring the temperature profiles in a BHE. Lee [27] has investigated 84 the effect of vertical heterogeneities of the soil thermal conductivity, concluding that the adoption 85 of depth-averaged thermal parameters is appropriate. Chung and Choi [28] have found that an 86 increase of the fluid flow rate reduces the heat transfer rate per unit length. Delaleux et al. [29] have 87 studied the increase of the thermal conductivity of grouts with the addition of graphite flakes, 88 concluding that a noticeable heat transfer improvement is achieved by BHEs. Jun et al. [30] have 89 evaluated the influence of running time, pipe spacing, grout conductivity, borehole depth, fluid flow 90 rate, inlet fluid temperature and soil type on the heat transfer length and on the thermal resistance of 91 borehole and soil. Michoupoulos and Kiriakis [31] have found a non-linear relation between the 92 BHE length and the heat pump consumption, which can be used for optimization processes in the 93 dimensioning of large plants. The aforementioned studies deal with single or few parameters, but a 94 thorough comparative analysis of all these factors together is still missing, and constitutes the 95 objective of this work. The functioning of a single BHE was simulated for 30 years, using a 96 benchmark cyclic thermal load and changing the operational parameters of the scenario. The 97 resulting fluid temperatures at the end of the BHE were processed and used to estimate the COP of 98 the heat pump and its annual energy consumption under different conditions. On the basis of the 99 results it is possible to draw some practical conclusions on the margins of improvement of BHEs 100 and on the proper choice of soil parameters for the simulations.

101

102 **2.** The modelling framework

103 The sensitivity analysis has been carried out on the design parameters of the BHE (geometrical 104 setting, properties of the materials, flow rate etc.) and on the physical properties of the soil and the

aquifer (thermal conductivity, groundwater flow velocity etc.), with the aim of evaluating their
relative impact on the performances of a GSHP (i.e. evolution of the heat carrier fluid temperatures,
energy consumption of the heat pump) in a realistic scenario and in long-term perspective.

The case study involves the simulations of the heating system of a house in the North of Italy, with a heated surface of 150 m² and a good thermal insulation. A geothermal heat pump connected to a BHE with a single U-pipe configuration is used only for heating. A cyclic thermal load (see Fig.2) has been set, with a total heat abstraction of 12 MWh per year (80 kWh m⁻²y⁻¹), which is equivalent to the energy produced by 1200 m³ of methane or 1250 l of gasoil using an efficient condensation boiler. The simulations last for 30 years, which is a sufficiently long time span to assess the longterm sustainability of the thermal exploitation of the soil.

115

The simulation of the heat exchange of the BHE with the soil and the aquifer system has been performed with FEFLOW 6.0, a 3D finite element flow and solute/heat transport model [32, 33] that includes specific tools for the simulation of Borehole Heat Exchangers [23, 24]. The software solves the coupled equations of flow and heat transport in the soil, and the BHE is modelled as an internal boundary condition of the 4th kind (thermal well).

121 The heat transport occurs by conduction (driven by thermal gradients), advection (due to the 122 groundwater flow) and dispersion (due to deviations from the average advective velocity), which 123 are described by the heat conservation equation in the porous medium:

124
$$\frac{\partial}{\partial t} \Big[\left(\varepsilon \rho_w c_w + (1 - \varepsilon) \rho_s c_s \right) T \Big] + \frac{\partial}{\partial x_i} \left(\rho_w c_w q_i T \right) + \frac{\partial}{\partial x_i} \Big[\left(\lambda_{ij}^{cond} + \lambda_{ij}^{disp} \right) \frac{\partial T}{\partial x_j} \Big] = H$$

125

126 where ε is the porosity, ρ_s and ρ_w are the density of the solid and liquid phase, c_s and c_w are the 127 specific heat of the solid and liquid phase, *T* is the temperature (which has been assumed equal for

both the phases), x_i is the i-th axis (i.e. $x_1 \equiv x, x_2 \equiv y, x_3 \equiv z$) and q_i is the i-th component of the Darcy velocity (i.e. relative to the i-th axis), and *H* is the heat source or sink (the BHE in this case), The first term of Eq.1 describes the soil temperature variation with time, involving the porosity ε and the heat capacity of the solid matrix $(\rho c)_s$ and of water $(\rho c)_w$.

132 The second term describes the advection, which depends on the Darcy velocity q.

133 The conduction and dispersion are respectively described by the tensors of the thermal conductivity 134 λ_{ij}^{cond} and λ_{ij}^{disp} (third term of Eq.1):

2

3

135
$$\lambda_{ij}^{cond} = \begin{cases} (1-\varepsilon)\lambda_s + \varepsilon\lambda_w & \text{for } i = j \\ 0 & \text{for } i \neq j \end{cases}$$

136

137
$$\lambda_{ij}^{disp} = \rho_w c_w \left[\alpha_T q \delta_{ij} + (\alpha_L - \alpha_T) \frac{q_i q_j}{q} \right]$$

138

139 where λ_s and λ_w are the thermal conductivities of the solid matrix and of groundwater, α_L and α_T 140 are the longitudinal and the transverse dispersivity (with respect to the direction of groundwater 141 flow) and q is the modulus of the Darcy velocity.

142

The temperature of the soil at the borehole wall, calculated by the 3D finite-element modelling 143 144 code, is used to solve the balance of the thermal fluxes inside the BHE according to the Thermal 145 Resistance and Capacity Model (TRCM) of Bauer et al. [34]. The BHE is decomposed into 146 different elements (inlet and outlet pipe, grout zones, borehole wall), which are represented by the 147 nodes of the circuit, connected by thermal resistances, which depend on the geometrical settings and 148 the physical properties of the materials. Thermal energy conservation equations are solved, which 149 describe the balance of thermal fluxes between the components of the BHE, and the temperature of 150 each component is calculated [23]. Since no abrupt changes occur in the thermal load, the analytical

method based on Eskilson and Claesson's solution [35], which considers a stationary equilibrium between the soil and the BHE, has been used in the simulations in order to reduce the computational time if compared to the Al Khoury et al.'s [36, 37] transient model.

154

A very large square mesh domain, with a side of 1000 m and a thickness of 150m, has been used to avoid boundary effects on the computed BHE fluid temperatures. The 31 flat slices are equally spaced (5m of distance) and the total number of nodes is 15531. The mesh density has been set using the "BHE node rule" [38], positioning the nodes around the BHE on the vertexes of a regular hexagon, with a radius of 0.46 m (6.13 times the borehole radius), since Diersch et al. [24] proved that this mesh density achieves a higher precision in the results, even when compared with finer meshes.

162

163 The thermal balance of the soil around the BHE has been reproduced choosing appropriate 164 boundary conditions. The temperature of the soil is almost constant through the year and, at an 165 infinite distance from the BHE, it is not affected by the thermal exchange. Constant temperature values (1st kind heat transport b.c.) have therefore been imposed at the lateral boundaries of the 166 167 domain, at least 500 m far away from the BHE. The heat flux coming from the deep layers of the Earth (geothermal flux), which has a mean value of 0.065 Wm^{-2} on the continental crust [39], 168 169 induces a temperature vertical gradient with typical values around 0.03 °C/m. According to these considerations, a temperature of 12°C has been set at the border or the first slice (which is a typical 170 171 value of the annual mean air temperature in Northern Italy), incrementing the temperatures of 0.15°C every 5m of depth (0.03 °C/m). The initial conditions have been set consistently with the 172 173 boundary conditions, with a homogeneous distribution of the soil temperature at each slice.

174

An unconfined aquifer, with a water table depth of 20m in the centre of the mesh (where the BHE is
positioned), has been modelled assigning constant hydraulic head (1st kind) flow boundary

177 conditions along the mesh borders. A homogeneous and isotropic hydraulic conductivity 178 $(K = 10^{-4} m/s)$ has been assigned, and different hydraulic gradients, ranging between 1‰ and 20‰, 179 have been imposed to change the groundwater flow velocity. Also different values of the saturated 180 thickness of the phreatic aquifer have been adopted, ranging from 10m to 50m in the middle of the 181 mesh, where the BHE is located.

182

A large set of simulations has been run in order to ascertain the influence of design parameters (length, pipe spacing, pipe diameter, heat carrier fluid and its flow rate, grout thermal conductivity), soil thermal (thermal conductivity of the solid matrix, thermal dispersivity) and hydrogeological properties (groundwater flow velocity, aquifer saturated thickness) on the performances of a BHE over a long operation period (30 years).

The adopted values of the BHE length range between 50 and 100 m, using a default value of 75 m. The borehole diameter is 0.15 m for all the simulations, and the HDPE pipes have an external diameter of 32mm and a wall thickness of 2.9 mm. The pipe spacing depends on the kind of spacers and from the pipe curvature given by the coil shape, which they keep even when they are unrolled: it varies therefore with depth and could not be known precisely. Different values have therefore been adopted, ranging from 35 to 117 mm between the pipe centres.

A set of simulation has been run to assess the performances of the most commonly adopted heat carrier fluids, and also different flow rates have been assigned $(0.1 \div 0.7 \text{ ls}^{-1} \text{ with propylene glycol at}$ 25% weight concentration). The default fluid is calcium chloride at 20% weight, which proved to be the most performing one.

The thermal conductivity of the BHE filling can vary in a wide range, and values between 1 and 5 Wm⁻¹K⁻¹ have therefore been adopted, while its thermal capacity does not experience great variations, and hence a unique value (2 MJm⁻³K⁻¹) has been used for all the simulations.

201 Some of the thermal and hydrogeological parameters of the soil have been kept constant for all the simulations, like the thermal properties of water ($\lambda_w = 0.6 Wm^{-1}K^{-1}$ and (ρc)_w = 4.2 MJm^{-3}), the 202 thermal capacity of the soil solid phase ((ρc)_w = 2.52 MJm⁻³) and both the total and the effective 203 porosity (respectively $\varepsilon = 0.3$ and $n_e = 0.2$), while the others have been changed to assess their 204 influence on the performances of the geothermal systems. As the heat transport occurs by 205 conduction, advection and dispersion, large ranges of the solid phase thermal conductivity (1÷3 206 $Wm^{-1}K^{-1}$), the Darcy velocity of groundwater (0÷17.32 md⁻¹) flow and the longitudinal/transverse 207 208 thermal dispersivity (0.5 m) have therefore been investigated.

209

210 The time series of the borehole fluid temperatures (Fig.3A) have been processed, calculating a 211 cumulative temperature distribution (Fig.3B) during the heating seasons over the whole simulation 212 period (30 years), which serves as a synthetic indicator to compare the different cases and to draw 213 conclusions on the energetic performance of the system. Observing the fluid temperature duration 214 curves in Fig.4 and Fig.5, one can understand how long will the heat pump work in a certain source 215 temperature range. For example, Fig.4A shows that, for a 75m long BHE, the mean fluid 216 temperature is below 0°C for the 19.51% of the heating period (say, 41.37 days a year), while this 217 percentage rises up to 50.86% for a 50m long borehole (107.83 days a year).

218

The Coefficient of Performance (COP), which is the ratio between the heating power delivered to the building and the electrical power absorbed by the heat pump, depends on the temperatures of the heat source (the BHE fluid) and of the heat sink (the heating terminals of the building). The relationship of COP from fluid temperatures has been approximated with a linear formula:

$$223 COP = a + b \cdot T_{t}$$

224

where T_f is the average fluid temperature between the inlet and outlet pipes of the BHE, while *a* and *b* depend on the heating terminal. For this study, we have set a = 4 and $b = 0.1 K^{-1}$, which are typical values for radiant panels at 35°C.

228 The estimated COP values at each time step (COP_i) have been used to calculate the energy 229 consumption of the heat pump:

230
$$HPC = \sum_{i=1}^{n} \frac{BHL_i}{COP_i} \cdot \Delta t$$

231

where BHL_i is the value of the BHE heat load at the i-th time step and Δt is the length of the constant time step (1 day). The electricity consumed by the heat pump gradually increases, as the soil and the BHE fluid is gradually cooling: the average value of yearly electricity consumption in the operation period (30 years) has been therefore used to evaluate the energy performance of the different BHE settings (Fig.6).

5

237

238 **3. Results and discussion**

The results of the long-term BHE simulations have been processed and compared in order to understand which is the relative importance of each parameter on the performances of the system and which is the margin of error due to the uncertainty in its determination, in particular for soil properties. Statistics about the calculated fluid temperatures (average, RMSE), the Seasonal Performance Factor (SPF) and the heat pump consumption for each simulation are summarized in the tables reported in the supporting information.

245

The length of the Borehole Heat Exchanger(s) plays a crucial role in the design process, because it accounts for about half of the total installation cost in single-house plants (see Blum et al. [40]). Varying the BHE length between 50 and 100m, we observe a strong variation of the cumulate 249 distributions of the average fluid temperatures (Fig.4A) and of the value of the minimum fluid 250 temperature, which is a critical parameter in the operation of a GSHP. The effect of the length 251 increase is non-linear and diminishes for larger BHE sizes: for example, incrementing the length by 252 between 50 and 75 m results in an increment of 2.80°C in the mean temperature, and of 1.58°C 253 when the increment is from 75 m to 100 m; the minimum inlet temperatures are incremented 254 respectively by 4.15°C and 1.92°C in the same ranges. The differences in the distributions of fluid 255 temperatures also have a noticeable impact on the energy expense of the heat pump, as shown in 256 Fig.6A. As for the cumulate distributions of the fluid temperatures, the effect of additional BHE length is reduced as the borehole depth increases (- 5.88% between 50 m and 75 m, -2.77% between 257 258 75 m and 100 m).

259

260 The improvement of the energy performance with longer exchangers is compensated by a rise in the 261 installation costs, which are the main drawback of geothermal heat pumps. In the dimensioning of 262 BHE fields, usually a minimum and/or maximum fluid temperature constraint is imposed, and the 263 minimum required borehole size is calculated [15, 16]. This approach minimizes the installation 264 costs, but the maintenance costs are not taken into account, and the extra-cost due to a low COP can 265 overcome the initial saving incurred with a smaller drilled depth. Starting from the results of the 266 sensitivity analysis on the length of the BHE, we have considered the typical electricity and BHE 267 installation costs of Italy (see Tab. 2) and calculated the total costs of installation and maintenance 268 of the GSHP over a lifetime of 30 years. Since the unit cost of electricity is likely to increase over 269 the next few decades, the analysis took into account different increase rates, in the range 270 between0% and 5%. In Fig.7, the ration between the lifetime cost for each BHE length and the most 271 expensive solution for each scenario of energy cost increase is shown, to identify the optimal size 272 for each case. We observe that higher increments of the unit cost of electricity enlarge the optimal 273 range of the BHE length, and shift it towards larger values; although it is not shown in the graph, a 274 decrease of the drilling cost also achieves the same effect. GSHPs need larger investments 275 compared to the other heating and cooling plants, and loan rates have been also considered when
276 evaluating the optimal length. Nevertheless, the influence of the interest rate on the total cost of the
277 plant over its lifetime proved to be negligible, compared to the cost of electricity and its increasing
278 trend.

279 A default length of 75 m was used in the other simulations, since it proved to be a reasonable choice 280 for most of the scenarios depicted in Fig.7. The considerations on BHE length that we have made 281 here concern only the lifetime cost of the plant, without taking into account the effects of very low 282 fluid temperatures. For example, if a GSHP operates at temperatures below 0°C for a sufficiently 283 long time, ground freezing can occur, and the borehole grouting can be fractured by freezing-284 thawing cycles. In addition, the viscosity of the heat carrier fluid increases as the temperature 285 decreases, therefore the energy consumption of the circulation pump also increases. A low 286 temperature threshold should therefore be established, which excludes some of the BHE lengths 287 considered in this analysis: for example, setting a minimum inlet temperature of -3°C excludes 288 lengths below 70 m.

289

290 Although the borehole depth exerts the greatest influence on the economic balance of a BHE 291 installation, there are also other factors which have to be taken into account. In the U-pipe BHEs 292 (both single and double), which are the most diffused kind of installation, the pipes should be put as 293 far as possible, to reduce both the thermal resistance of the exchanger and the heat exchange 294 between the inlet and the outlet pipes (thermal short-circuit), which impair the performances of 295 these systems. The thermal conductivity of the borehole filling plays an important role: a higher 296 value reduces the borehole resistivity, but also the grout-to-grout resistance, which prevents the 297 thermal short-circuit. Both these factors have been taken into account in the simulations, according 298 to the borehole resistance model of Bauer et al. [34]. The distance between the pipe centres has 299 been varied between 35 mm (i.e. 3 mm between the pipe walls) and 117 mm (i.e. 0.5 mm between 300 the pipe wall and the borehole wall), and the thermal conductivity of the grout has been varied

between 1 $\text{Wm}^{-1}\text{K}^{-1}$ (i.e. a poor grout) and 5 $\text{Wm}^{-1}\text{K}^{-1}$ (i.e. special grouts with highly conductive 301 302 graphite flakes [29]). Usually, the grouts employed for BHEs have a thermal conductivity of 2÷2.5 Wm⁻¹K⁻¹, but this value can dramatically decrease due to an incorrect mixing, an excessive water 303 content or an insufficient concentration of thermal additives [41]. Observing the cumulate 304 305 distributions of the fluid temperatures (Fig.4B-C), we understand that the influence of the thermal 306 conductivity of the grout is very large when the pipe spacing is reduced; on the other hand, a grout 307 with a high thermal conductivity can compensate the negative effects of an insufficient pipe spacing 308 on both the minimum fluid temperatures and the energy consumption of the system (Fig.6B). For example, if a common geothermal grout is used $(\lambda_g = 2Wm^{-1}K^{-1})$, the consumption of the heat 309 pump diminishes of the 1.99% as the pipe distance is increased from 35mm to 117 mm; on the other 310 hand, if a highly conductive grout ($\lambda_g = 5Wm^{-1}K^{-1}$) is used, this difference is reduced to the 0.64%, 311 meaning that special grouts noticeably reduce the effect of an insufficient pipe spacing. 312

313

314 The fluid circulated into the closed pipe loop is usually a mixture of water and antifreeze. The flow 315 rate and the physical properties of this fluid (viscosity, thermal capacity, thermal conductivity) 316 influence the borehole thermal resistance [42]. The main drawbacks of increasing the concentration 317 of the antifreeze additive are a noticeable increase of viscosity, a slight decrease of the thermal 318 conductivity and an additional cost (say $2 \div 4 \notin /1$, depending on the kind of ethanol or glycol); in 319 addition, the antifreeze is a potential source of contamination in case of a pipe leak, and the anti-320 corrosion additives can inhibit the bacterial degradation [43]. All these adverse side effects should 321 be minimized when choosing the anti-freeze additive. Simulations have been carried out 322 considering the most common anti-freeze mixtures: propylene glycol (PG) at 25% and 33% volume concentration, ethanol (ETH) at 24% vol., calcium chloride (CaCl₂) at 20% weight concentration. 323 324 Their physical properties are reported in Tab. 1, where also the boundaries of the laminar and of the turbulent regime are shown, since the thermal resistance is much smaller in turbulent one [42]. The 325

default flow rate is 0.5 ls⁻¹, which is a typical value for GSHPs. The results (Fig.4D and Fig.6C) show that calcium chloride solutions permit to achieve an appraisable gain in the energy performance (compared to PG25%, minimum temperature: +2.94°C; heat pump consumption: -4.01%), due to their smaller viscosity and their higher thermal conductivity; in addition, it is much cheaper than the other antifreeze additives. On the other hand, the use of saline solutions as a heat carrier fluid requires the adoption of specific anti-corrosion components.

The other antifreeze mixtures show negligible variations of the fluid temperatures and of energetic performances. As the thermal resistance diminishes when higher flow rates are circulated, seven simulations (fluid: PG25%, flow rates: $0.1\div0.7 \text{ ls}^{-1}$) have been run to quantify its contribution for a better efficiency of the GSHP. We observe that the energy consumption of the heat pump is reduced of the 4.4% between 0.1 and 0.7 ls⁻¹; nevertheless, circulating larger flow rates implies also a higher energy expense for the circulation pump. We have therefore quantified the distributed friction losses along the 75m long using the explicit approximation of the Prandtl formula (Eq.6) for smooth pipes:

339
$$\lambda_0 = \frac{0.25}{\left[\log_{10}\left(\frac{5.7}{\text{Re}^{0.9}}\right)\right]^2}$$

340

341 where $\lambda_0 = 2g \frac{d_{ip}}{u^2} \cdot J$ is the non dimensional friction loss, d_{ip} is the pipe internal diameter, g is the 342 gravity acceleration, J is the hydraulic gradient in the pipes. 6

7

343 The energy consumption of the circulation pump increases rapidly with the fluid flow rate (Q_f) :

344
$$CPC = \frac{J \cdot 2L \cdot \rho_f \cdot g \cdot Q_f}{\eta} \cdot t_{func} = \frac{16 \cdot \lambda_0 \cdot L \cdot \rho_f \cdot g}{\eta \cdot \pi^2 \cdot D^4} Q_f^{-3} \cdot t_{func}$$

345

where ρ_f is the density of the heat carrier fluid, *L* is the BHE length [L] and t_{func} is the operation time per year. An energy yield $\eta = 0.8$ has been assumed for the calculation of *CPC*. Fig.8 shows the strong impact of the flow rate on the total energy consumption (circulation and heat pump). In particular, a strong variation occurs when switching from laminar to transition regime (between 0.2 and 0.3 l/s), with a reduction of 2.07% for the total energy consumption, while the minimum values lie in a range of flow rates (for this case, $0.3\div0.5$ ls⁻¹). Noticeable differences are observed in the minimum temperature, meaning that higher flow rates can be adopted when larger amounts of heat are extracted from the soil, in order to avoid the freezing of the ground, or to reduce its extent.

355

While the design parameters can be determined with an acceptable precision, the real issue of GSHP modelling is the knowledge of the physical parameters of the soil. The heat transport around the BHE is mainly conductive, especially if no significant groundwater flow occurs, therefore the most important soil physical parameter is the thermal conductivity of the porous medium λ_{ij}^{cond} (see Eq.2).

The thermal conductivity of the solid matrix (λ_{e}) is the parameter which can vary in the widest 361 362 range, depending on the lithology, the grain size, the water saturation etc.. A wide range of values has been explored in the simulations $(1 \div 3 \text{ Wm}^{-1} \text{K}^{-1})$, and the graphs of the cumulate distribution of 363 the fluid temperatures (Fig.5E) and of the heat pump energy consumption (Fig.6D) show that 364 365 thermal conductivity has a very strong influence on the performances of the system, compared to the BHE length. Especially in smaller installations, this parameter is not measured in situ, but low-366 367 precision data from literature are adopted (e.g. the German norm VDI 4640 [44]). For example, the thermal conductivity of a moraine ranges between 1 and 2.5 Wm⁻¹K⁻¹, for which we observe a 368 369 difference of 5.66°C in the minimum temperature, and 12.5% in the power consumed by the heat 370 pump. An imprecise knowledge of this parameter results therefore in a strong uncertainty in the 371 simulation of the plant, which has to be overcome e.g. with a Thermal Response Test [45].

373 The presence of a subsurface flow has been proved to be beneficial for the performances of closed-374 loop geothermal heat pumps. Indeed, groundwater flow activates advection and thermal dispersion, enhancing the heat transport around the BHE and spreading the thermal disturbance further away. 375 376 Chiasson et al. [25] demonstrated that the advection has a considerable impact only in coarse-377 grained soil (sands and gravels) and in fractured aquifers (e.g. karst limestone), while Wang et al. 378 [26] stressed the importance of the saturated thickness, which can vary through the year, influencing 379 also the results of Thermal Response Tests [12]. A set of simulations with different flow velocities 380 and saturated thicknesses has been run therefore to quantify the positive effect of groundwater flow in a typical sand aquifer ($K = 10^{-4} m/s$). 381

As shown in Fig.5B-C and Fig.6E, the influence of the Darcy velocity on the performances of the system is much stronger than the variation induced by different saturated thicknesses. This means that the contribution of the advection can be taken into account, but precise values are needed to avoid undersized design; on the other hand, variations in the saturated thickness - e.g. due to seasonal level variations in surface water bodies - do not exert a strong influence on the operation of GSHPs, if the gradient does not experience significant variations.

388

389 When modelling heat transport in an aquifer, one should consider also the dispersion, which is a 390 strong mechanism of heat transport. The thermal dispersivity has been considered as a scale-391 dependent parameter, as reported in literature [46]. Sethi and Di Molfetta [21] adopted $\alpha_L = 10 m$ and $\alpha_T = 1 m$ for the heat transport simulation around a municipal solid waste landfill. Erol [47] 392 assumed $\alpha_L = 2 m$ and $\alpha_T = 0.2 m$ for the simulation of a 100 m long BHE. Molina-Giraldo et al. 393 394 [48] analyzed the extension of the thermal plume downstream of a BHE, for different values of groundwater flow Darcy velocity ($q = 10^{-8} \div 10^{-5} m/s$) and for different values of thermal 395 dispersivity ($\alpha_L = 0 \div 2m$), discovering that thermal dispersion reduces the extent of a reference 396 397 isotherm (e.g. $+1^{\circ}$ C) of the deviation from the undisturbed soil temperature.

Wagner et al. [49] also analyzed the effect of α_L for Thermal Response Tests in presence of groundwater flow, concluding that thermal dispersion can lead to a strong overestimation of the thermal conductivity of the soil. This is confirmed by the cumulate distribution of the average fluid temperatures for a Darcy velocity of 4.32 m/day (Fig.5D), which prove that the thermal dispersion is a great factor of uncertainty when modelling BHE fields in presence of subsurface flow. A rule of thumb that is usually employed in the solute transport [50] is:

8

$$\alpha_L = 0.1 L_p$$

405

404

where L_p is the spatial scale of the dispersion phenomenon. The concept of "scale" is not 406 univocally defined for GSHPs: using the BHE diameter (i.e. $\alpha_L = 0.1m$ or less) or its length (i.e. 407 $\alpha_L = 10 m$) would imply a difference of some 8÷10°C for the minimum fluid temperature and more 408 409 than 15% for the electricity consumption of the heat pump (see Fig.6F). It is therefore advised not 410 to rely on thermal dispersion when designing BHE fields, until field tests will be carried to estimate 411 the thermal dispersivity in real-scale setups: especially if a thick and conductive aquifer is present, 412 the overestimation of the thermal dispersivity would lead to an under-dimensioning of the GSHP 413 with a detrimental effect on its long-term sustainability.

414

415 **4.** Conclusions

In this work, the most important parameters which influence the performances of Ground Source Heat Pumps have been thoroughly analyzed, running long-term simulations and estimating the energy consumption of the heat pump for each setting. Most of these factors have been already analyzed in other works, but none of them considered all the parameters together, using the same modelling framework and considering the effect on the lifetime of a GSHP. The analysis of the BHE design parameters (length, pipe spacing, fluid, grout) permits to understand which are the margins of improvement, while the physical parameters of the soil (thermal conductivity and 423 dispersivity, groundwater flow) have been analyzed in order to understand their effect on the424 uncertainty in the project phase.

The results of the simulations prove that the length of the BHE is the most important parameter in the design of a GSHP. Indeed, increasing the borehole depth results in a reduction of the thermal disturbance in the subsoil and therefore to achieve a higher efficiency of the heat pump, but also a larger investment is needed for the installation.

429 An optimum length should be found, which minimizes the total cost over the plant lifetime, 430 considering also the trend of increase of the unit cost of electricity. While the drilled depth has an 431 appraisable impact on the initial investment, there are also other important factors to be considered 432 for the optimization of BHEs, like the pipe arrangement, the grout and the heat carrier fluid. A large 433 pipe spacing and a highly conductive grout, reducing the heat losses in the heat exchange with the 434 soil, achieves an appraisable reduction of the energy costs for the heat pump with a negligible 435 expense, compared to the borehole drilling. For the circulation pump, a trade-off can be found for 436 the choice of the correct flow rate for the heat carrier fluid, allowing the minimization of both the 437 energy losses due to the thermal resistance and the friction losses due to the circulation of the fluid. 438 The antifreeze and its concentration heavily influence the energy performance of GSHPs, in 439 particular the borehole resistance and the power consumed by the auxiliary plants. The saline 440 solutions, with a smaller viscosity compared to ethanol and glycols, permit to reduce all these 441 energy losses, although special components are needed to avoid corrosion problems. Optimizing the 442 design and the installation of BHEs is useless without a thorough characterization of the subsoil, 443 which has a large influence on the performances of these systems. When no groundwater flow 444 occurs, the thermal conductivity is the most important parameter for the dimensioning of BHEs. 445 The technical literature provides wide ranges of the thermal conductivity for each lithology, which 446 can vary due to porosity, saturation and other factors; in-situ Thermal Response Tests are therefore 447 strongly advised for large plants to avoid under or over dimensioning. The advection enhances the 448 performances of GSHP, and the groundwater flow should be taken into account using conservative 449 values of hydraulic conductivity and gradient, unless they are known by field tests. On the other 450 hand, it is risky to consider also the beneficial effect of heat dispersion, because the thermal 451 dispersivity is still scarcely known in real-scale BHEs. In situ tests to estimate these parameters 452 would be highly desirable to simulate the behaviour of BHE fields with a better precision.

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567 Nomenclature

568	BHL	Total annual BHE Heat Load (kWh y ⁻¹)
569	BHL_i	BHE Heat Load at the i-th time step (kW)
570	C_{f}	Groundwater specific heat (J kg ⁻¹ K ⁻¹)
571	C _s	Aquifer solid matrix specific heat (J kg ⁻¹ K ⁻¹)
572	СОР	Coefficient of Performance of the heat pump (dimensionless)
573	CPC	Circulating Pump Consumption (kWh y ⁻¹)
574	$d_{_{ip}}$	Internal pipe diameter (m)
575	$d_{_{op}}$	External pipe diameter (m)
576	g	Gravity acceleration (m s ⁻²)
577	Н	Heat source/sink (W/m ³)
578	HPC	Total annual Heat pump energy consumption (kW y ⁻¹)
579	HPC_i	Power consumed by the heat pump at the i-th time step (kW)
580	$-\frac{\partial h}{\partial x}$	Hydraulic gradient in the aquifer (dimensionless)
581	J	Hydraulic gradient in the BHE pipes (dimensionless)
582	K	Hydraulic conductivity of the aquifer (m s^{-1})
583	L	Length of the BHE (m)
584	L_p	Scale dimension (m)
585	n _e	Effective porosity or specific yield of the aquifer (dimensionless)
586	\mathcal{Q}_{f}	Flow rate of the heat carrier fluid (1 s^{-1})
587	q	Darcy velocity of groundwater flow (m s ⁻¹)
588	q_i	i-th component of the Darcy velocity (m s^{-1})

589	Re	Reynolds number (dimensionless)
590	RMSE	Root Mean Square Error
591	Т	Temperature of the soil, both solid and fluid phase (°C)
592	T_{f}	Average fluid temperature (°C)
593	t _{func}	Functioning time of the circulation pump (d y^{-1})
594	T _{in}	Inlet fluid temperature (°C)
595	T _{out}	Outlet fluid temperature (°C)
596	T_s	Soil temperature at the borehole interface (°C)
597	и	Flow velocity in the BHE pipes (m s ⁻¹)
598	w	Distance between the centres of the pipes in a BHE (m)
599	Greek letters	5
600		
000	$lpha_{\scriptscriptstyle L}$	Longitudinal thermal dispersivity (m)
601	$lpha_{\scriptscriptstyle L}$ $lpha_{\scriptscriptstyle T}$	Transverse thermal dispersivity (m)
601 602	$lpha_{\scriptscriptstyle L}$ $lpha_{\scriptscriptstyle T}$ $arepsilon$	Transverse thermal dispersivity (m) Porosity of the soil (dimensionless)
600601602603	$lpha_{\scriptscriptstyle L}$ $lpha_{\scriptscriptstyle T}$ $arepsilon$ η	Longitudinal thermal dispersivity (m)Transverse thermal dispersivity (m)Porosity of the soil (dimensionless)Energy yield (dimensionless)
600601602603604	$egin{array}{c} lpha_{_L} \ lpha_{_T} \ arepsilon \ $	Longitudinal thermal dispersivity (m) Transverse thermal dispersivity (m) Porosity of the soil (dimensionless) Energy yield (dimensionless) Non-dimensional friction loss (dimensionless)
 600 601 602 603 604 605 	$egin{array}{llllllllllllllllllllllllllllllllllll$	Longitudinal thermal dispersivity (m) Transverse thermal dispersivity (m) Porosity of the soil (dimensionless) Energy yield (dimensionless) Non-dimensional friction loss (dimensionless) Thermal conductivity of the heat carrier fluid (W m ⁻¹ K ⁻¹)
 600 601 602 603 604 605 606 	$egin{array}{c} lpha_{L} \ lpha_{T} \ arepsilon \ a$	Longitudinal thermal dispersivity (m) Transverse thermal dispersivity (m) Porosity of the soil (dimensionless) Energy yield (dimensionless) Non-dimensional friction loss (dimensionless) Thermal conductivity of the heat carrier fluid (W m ⁻¹ K ⁻¹) Thermal conductivity of the grout (W m ⁻¹ K ⁻¹)
 600 601 602 603 604 605 606 607 	$egin{array}{c} lpha_{L} \ lpha_{T} \ arepsilon \ arepsilon \ arepsilon \ arphi_{T} \ arepsilon \ arphi_{T} \ arepsilon \ arphi_{T} \ arphi_{0} \ arphi_{f} \ arphi_{g} \ arphi_{g} \ arphi_{p} \ a$	Longitudinal thermal dispersivity (m) Transverse thermal dispersivity (m) Porosity of the soil (dimensionless) Energy yield (dimensionless) Non-dimensional friction loss (dimensionless) Thermal conductivity of the heat carrier fluid (W m ⁻¹ K ⁻¹) Thermal conductivity of the grout (W m ⁻¹ K ⁻¹) Thermal conductivity of the BHE pipes (W m ⁻¹ K ⁻¹)
 600 601 602 603 604 605 606 607 608 	$egin{array}{c} lpha_{L} \ lpha_{T} \ arepsilon \ a$	Longitudinal thermal dispersivity (m) Transverse thermal dispersivity (m) Porosity of the soil (dimensionless) Energy yield (dimensionless) Non-dimensional friction loss (dimensionless) Thermal conductivity of the heat carrier fluid (W m ⁻¹ K ⁻¹) Thermal conductivity of the grout (W m ⁻¹ K ⁻¹) Thermal conductivity of the BHE pipes (W m ⁻¹ K ⁻¹) Thermal conductivity of the solid matrix of the soil (W m ⁻¹ K ⁻¹)
 600 601 602 603 604 605 606 607 608 609 	$egin{array}{llllllllllllllllllllllllllllllllllll$	Longitudinal thermal dispersivity (m) Transverse thermal dispersivity (m) Porosity of the soil (dimensionless) Energy yield (dimensionless) Non-dimensional friction loss (dimensionless) Thermal conductivity of the heat carrier fluid (W m ⁻¹ K ⁻¹) Thermal conductivity of the grout (W m ⁻¹ K ⁻¹) Thermal conductivity of the BHE pipes (W m ⁻¹ K ⁻¹) Thermal conductivity of the solid matrix of the soil (W m ⁻¹ K ⁻¹) Groundwater thermal conductivity (W m ⁻¹ K ⁻¹)

611	λ_{ij}^{disp}	Thermal conductivity for dispersion (W $m^{-1} K^{-1}$)
612	$oldsymbol{ ho}_{f}$	Density of the heat carrier fluid (Kg m ⁻³)
613	$ ho_{s}$	Density of the solid matrix of the soil (Kg m ⁻³)
614	$ ho_{_w}$	Density of groundwater (Kg m ⁻³)
615	$\left(ho c ight)_{_{f}}$	Thermal capacity of the heat carrier fluid (J $m^{-3} K^{-1}$)
616	$\left(ho c ight)_{g}$	Thermal capacity of the grout (J m ⁻³ K ⁻¹)
617	$(\rho c)_s$	Thermal capacity of the solid matrix of the soil (J $m^{-3} K^{-1}$)
618	$(ho c)_{_{\scriptscriptstyle W}}$	Thermal capacity of the solid matrix of the soil $(J m^{-3} K^{-1})$

619 Tables

Fluid	T _{freezing} [°C]	λ _f [Wm ⁻¹ K ⁻¹]	С _f [Jkg ⁻¹ K ⁻¹]	ρ _f [kgm ⁻³]	μ _f [mPas]	Q_{lam} [ls ⁻¹]	Q_{turb} [ls ⁻¹]
Prop.Glycol 25%	-10	0.45	3974	1026	5.51	0.252	1.097
Ethanol 24.4%	-15	0.426	4288	972	5.85	0.283	1.229
Prop.Glycol 33%	-15	0.416	3899	1015	8.17	0.378	1.644
CaCl ₂ 20%	-20	0.54	3030	1186	4	0.158	0.689

620

621 Tab. 1 – Physical properties of the anti-freeze solutions used in the simulations: solidification temperature 622 $(T_{freezing})$, thermal conductivity (λ_f) , specific heat (c_f) , density (ρ_f) , dynamic viscosity (μ_f) , upper boundary 623 flow rate for the laminar regime (Q_{lam}) and lower boundary flow rate for the turbulent regime (Q_{turb}) .

624

Parameter	Values
6 kW heat pump + installation	6000€
BHE drilling + installation	70 €/m
Unit cost of electricity	0.22 €/kWh
Increment of the unit cost of electricity	0%, 1%, 3%, 5%

625

626 Tab. 2 – Installation and energy costs used for the optimization procedure of the BHE length.

627

628 Figure captions

Fig. 1 – Scheme of a Ground Source Heat Pump (GSHP): the Borehole Heat Exchanger (BHE) exchanges heat between the surrounding soil and the heat pump. A thermal storage tank reduces the frequency of start-up and stop of the heat pump. Radiant panels and fan coils are the most diffused heating terminals for GSHPs. If present, groundwater flow enhances the heat transport around the BHE, permitting to achieve better energy performances.

635 Fig. 2 – Building Heat Load (BHL) adopted as a benchmark for the BHE in the simulations.

636

Fig. 3 – A: Time series of the average fluid temperatures, detail of 5 years of simulation. B: Cumulate
distribution of the average fluid temperatures in the heating seasons.

639

Fig. 4 – Cumulate distributions of the average fluid temperatures for different values of BHE length (A), pipe
spacing (B), thermal conductivity of the grout (C) and heat carrier fluids (D).

642

Fig. 5 – Cumulate distributions of the average fluid temperatures for different values of the thermal conductivity
of the solid matrix of the soil (A), groundwater flow Darcy velocity with no thermal dispersion (B), Darcy
velocity and saturated thickness (C), thermal dispersivity (D).

646

Fig. 6 – Estimated annual heat pump energy consumption for different values of BHE length (A), pipe spacing
and grout conductivity (B), heat carrier fluids (C), solid-phase soil thermal conductivity (D), groundwater flow
Darcy velocity and saturated thickness (E) and thermal dispersivity (F).

650

Fig. 7 – Relative variation of the total cost of a GSHP over a lifetime of 30 years, for different BHE lengths
(50÷100m) and different increment rates of the unit cost of electricity (0÷5%).

653

Fig. 8 – Cumulate distributions of the average fluid temperatures (A) and electric power consumption of the heat
pump and circulation pump (B) for different fluid flow rates.

Supporting information

Tab. 1 – Summary of the results of the simulations: mean and RMSE of the average fluid temperature (T_f) , minimum values of the inlet temperature (T_{in}) , Seasonal Performance Factor (SPF) and annual Heat Pump Consumption (HPC).

Parameter	value	Parameter	value	mean (<i>T_f</i>) [°C]	RMSE (T_f) [°C]	min (T _{in}) [°C]	SPF [-]	HPC [kWh/y]
		BHE length [m]	50	1.83	5.69	-6.39	4.22	2845.8
			55	2.52	5.36	-5.34	4.29	2799.5
			60	3.07	5.01	-4.40	4.34	2765.6
			65	3.46	4.79	-3.94	4.37	2743.0
			70	4.14	4.54	-2.99	4.44	2704.0
			75	4.63	4.25	-2.24	4.48	2678.4
			80	4.99	4.06	-1.71	4.51	2661.4
			85	5.26	3.92	-1.49	4.53	2648.5
			90	5.62	3.76	-1.02	4.56	2631.3
			95	6.03	3.57	-0.47	4.60	2611.1
			100	6.21	3.46	-0.33	4.61	2604.2
pipe spacing [mm]	pipe spacing [mm] 35 grout th		1	0.34	7.16	-10.02	4.06	2956.2
	35		2	3.88	4.85	-3.74	4.42	2715.9
	35		3	4.18	4.59	-3.22	4.44	2701.2
	35		5	4.65	4.27	-2.27	4.48	2678.4
	55		1	0.62	6.92	-9.43	4.09	2931.5

Parameter	value	Parameter	value	$ \begin{array}{c} \operatorname{mean}\left(T_{f}\right) \\ \left[^{\circ}\mathbf{C}\right] \end{array} $	RMSE (T_f) [°C]	min (T _{in}) [°C]	SPF [-]	HPC [kWh/y]
	55		2	4.05	4.74	-3.46	4.43	2706.4
	55		3	4.29	4.49	-3.04	4.45	2695.5
	55		5	4.64	4.25	-2.20	4.48	2678.0
	80		1	1.32	6.39	-7.95	4.17	2877.8
	80		2	4.20	4.55	-3.10	4.45	2699.6
	80		3	4.60	4.27	-2.31	4.48	2680.9
	80		5	4.68	4.17	-2.06	4.48	2677.7
	100		1	4.23	4.77	-3.41	4.45	2695.1
	100		2	4.74	4.33	-2.33	4.49	2670.7
	100		3	4.67	4.18	-2.09	4.48	2677.6
	100		5	4.74	4.05	-1.97	4.48	2676.9
	117		1	3.09	4.95	-4.29	4.34	2766.9
	117		2	5.03	3.90	-1.36	4.51	2661.9
	117		3	5.07	3.86	-1.31	4.51	2660.7
	117		5	5.07	3.84	-1.32	4.51	2661.3
		heat carrier fluid	PG 33%	1.95	5.96	-6.74	4.23	2834.8
			ET 24%	1.98	5.79	-6.49	4.23	2835.0
			PG 25%	2.49	5.50	-5.62	4.29	2799.7
			CaCl ₂ 20%	4.43	4.43	-2.68	4.47	2687.4
		solid matrix thermal conductivity [Wm ⁻¹ K ⁻¹]	1	-0.63	6.08	-9.36	3.91	3072.7
			1.5	1.32	5.47	-6.86	4.14	2895.8
			2	2.73	4.99	-4.72	4.30	2790.5

Parameter	value	Parameter	value	$ \begin{array}{c} \mathbf{mean} \left(T_{f} \right) \\ [^{\circ}\mathbf{C}] \end{array} $	RMSE (T_f) [°C]	min (T _{in}) [°C]	SPF [-]	HPC [kWh/y]
			2.5	3.60	4.64	-3.70	4.38	2737.5
			3	4.63	4.25	-2.24	4.48	2678.4
saturated thickness [m]	55	groundwater flow Darcy velocity [m/d]	0.864	5.28	4.37	-1.83	4.54	2640.8
			1.728	5.99	4.37	-1.39	4.61	2601.0
			1.32	6.85	4.17	-0.33	4.70	2552.5
			8.64	7.53	3.77	0.80	4.76	2520.7
			17.28	8.21	3.29	2.14	4.82	2491.9
groundwater flow Darcy velocity [m/d]	0.864	saturated thickness [m]	10	4.67	4.32	-2.27	4.48	2676.2
			20	4.91	4.31	-2.06	4.51	2662.4
			50	5.25	4.36	-1.83	4.54	2642.6
groundwater flow Darcy velocity [m/d]	8.64	saturated thickness [m]	10	5.73	4.23	-1.20	4.59	2615.7
			20	6.43	4.07	-0.54	4.66	2576.8
			50	7.41	3.82	0.62	4.75	2526.6
groundwater flow Darcy velocity [m/d]	0.864	longitudinal thermal dispersivity [m]	0.1	5.32	4.35	-1.82	4.55	2639.0
			0.2	5.35	4.33	-1.81	4.55	2637.7
			0.5	5.46	4.27	-1.74	4.56	2632.0
			1	5.61	4.20	-1.35	4.57	2624.1
			2	5.72	4.11	-1.26	4.58	2619.1
			5	6.32	3.78	-0.44	4.63	2589.9
groundwater flow Darcy velocity [m/d]	1.728	longitudinal thermal dispersivity [m]	0.1	6.04	4.34	-1.32	4.62	2598.7
			0.2	6.10	4.30	-1.27	4.62	2596.1
			0.5	6.23	4.20	-0.89	4.63	2589.8

Parameter	value	Parameter	value	mean (T _f) [°C]	RMSE (T_f) [°C]	min (T _{in}) [°C]	SPF [-]	HPC [kWh/y]
			1	6.43	4.08	-0.80	4.65	2580.5
			2	6.62	3.91	-0.37	4.67	2572.3
			5	7.48	3.39	1.16	4.74	2534.0
groundwater flow Darcy velocity [m/d]	4.32	longitudinal thermal dispersivity [m]	0.1	6.91	4.14	-0.29	4.71	2550.0
			0.2	6.98	4.09	-0.26	4.71	2547.1
			0.5	7.26	3.91	0.29	4.73	2535.4
			1	7.60	3.68	0.93	4.76	2521.6
			2	8.09	3.33	1.86	4.80	2501.3
			5	8.94	2.78	3.44	4.86	2467.8

Tab. 2 – Summary of the results of the simulations with different values of heat carrier fluid flow rate (Q_f): mean and RMSE of the average fluid temperature (T_f), minimum values of the inlet temperature (T_{in}), Seasonal Performance Factor (SPF), annual heat pump consumption (HPC), circulating pump energy consumption (CPC) and total energy consumption (HPC+CPC).

Q_f [l/s]	$ \begin{array}{c} {\rm mean} \\ (T_f) [^{\circ}{\rm C}] \end{array} $	RMSE (T_f) [°C]	min (<i>T_{in}</i>) [°C]	SPF	HPC [kWh/y]	CPC [kWh/y]	HPC+CPC [kWh/y]
0.1	-0.52	7.34	-16.57	4.09	2932.4	1.3	2933.8
0.2	0.19	7.02	-11.68	4.13	2906.6	8.2	2914.8
0.3	1.71	5.96	-7.51	4.24	2830.6	24.0	2854.6
0.4	2.49	5.50	-5.62	4.29	2799.7	51.9	2851.6
0.5	2.76	5.24	-4.97	4.28	2800.9	94.5	2895.4
0.6	3.14	4.99	-4.17	4.30	2790.8	154.5	2945.3
0.7	3.14	4.94	-3.75	4.28	2803.1	234.4	3037.5



Fig.1







A



PG 33%

ETH 24%

100

PG 25% CaCl₂ 20%

80





q=4.32 m/d

q=8.64 m/d

q=17.28 m/d

100

80

















